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3 messages

Chi-Wang Li <onbehalfof@manuscriptcentral.com> Reply-To: chiwangli@gmail.com To: muh.sarkowi@eng.unila.ac.id, sarkov323@yahoo.com Tue, Aug 3, 2021 at 1:40 PM

03-Aug-2021

Dear Mr. Sarkowi:

It is a pleasure to accept your manuscript entitled "Reservoir Identification of Bac-Man Geothermal Field Based on Gravity Anomaly Analysis and Modeling" in its current form for publication in the Journal of Applied Science and Engineering. The comments of the reviewer(s) who reviewed your manuscript are included at the foot of this letter.

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Reviewer: 1

Comments to the Author

I congratulate the authors for their decision to improve their manuscript carefully and resubmit the revised version. The manuscript has been noticeably improved. I recommend the authors to choose stronger visualization style in their future works. Adding complementary information of any kind (in your study is e.g. thermal data) to the concluding plots (your cross sections), always, helps readers to understand and believe on the accuracy of the results.

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Comments to the Author The revised paper has answered the reviewer's questions. It is accetable for publication in JASE.

Muh Sarkowi <muh.sarkowi@eng.unila.ac.id> To: chiwangli@gmail.com Wed, Aug 4, 2021 at 7:18 AM

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Best regards Muh Sarkowi

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Chi-Wang Li <chiwangli@gmail.com> To: Muh Sarkowi <muh.sarkowi@eng.unila.ac.id>

Rahmat,

Great.

Best regards,

Professor Chi-Wang Li, Ph.D. Department of Water Resources & Environmental Engineering Wed, Aug 4, 2021 at 7:24 AM

Tamkang University, Taiwan Editor-in-Chief Journal of Applied Science and Engineering http://jase.tku.edu.tw/

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1	Reservoir Identification of Bac-Man Geothermal Field Based on
2	Gravity Anomaly Analysis and Modeling
3	Muh Sarkowi ¹ *, Rahmat Catur Wibowo ²
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5	35145, Lampung, Indonesia, *email: <u>muh.sarkowi@eng.unila.ac.id</u>
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8	Abstract
9	The Bac-Man geothermal field is located between the provinces of Albay and Sorsogon
10	on the Bicol Peninsula. Several geophysical modeling methods have been carried out to
11	identify the Bac-Man field geothermal system. This research focuses on 3D modeling and
12	analysis of gravity data which aims to identify the geothermal reservoir in the Bac-Man field.
13	This research includes several things, such as: spectrum analysis and the separation of
14	Bouger's anomalies; gradient analysis; and anomaly modeling. Based on the modeling results,
15	there are 3 low anomaly closures in the middle which are separated by high anomalies which
16	can be interpreted that the geothermal reservoir in the Bac-Man field may be divided into 3
17	reservoirs, namely the southern part (Cawayan and Tanawon sector), the eastern part (Boton
18	sector). and the northern part (Palayan - Inang Maharang sector) where each reservoir area is
19	separated by the presence of a fault or intrusion structure.

21 Keywords: Bac-Man field, Gravity, Modeling, Reservoir, Geothermal

23 1. Introduction

The Bacon - Manito (Bac-Man) geothermal field is located in the Podcol mountains on the island of Luzon, about 350 km southeast of Manila. The exploration of the Bac-Man geothermal field has been carried out since 1977. The reservoir model has an area of 23 km², a depth of 1500 m and a temperature of 240°C to 320°C. Based on geological, geochemical and geophysical data, a conceptual model of the Bac-Man geothermal system has been made (Austria, 2008).

Tugawin *et al.* (2015), carried out 2D Magnetotellurics inversion modeling in the Bac-Man field which shows the existence of 3 (three) geothermal reservoir areas, namely Palayan Bayan, Tikolob, and Malobago areas. These results are in accordance with the results of data interpretation of resisitivity vertical electric sounding (VES) and Schlumberger Resistivity Traversing (SRT). The Reservoir in the Palayan Bayan area has been in production since 1993, while the Tikolob prospect is on the west side of the Bac-Man field, separate from the Palayan Bayan system (Tugawin *et al.*, 2015).

Espartinez and See (2005) have conducted geochemical research in the Bac-Man Geothermal field to determine changes in reservoirs from their chemical properties. This is done to maintain the continuity of steam production in the Bac-Man field considering that of the 3 geothermal reservoir sectors, the Palayan Bayan and Cawayan sectors are still producing while the Boton sector has not been producing since 2009 (See, Fragata and Solis, 2005).

In this research, gravity data processing and modeling will be carried out to determine the geothermal reservoir in the Bac-Man field (Palayan Bayan, Cawayan and Boton). The results of the subsurface model derived from anomaly gravity analysis and the density distribution model resulting from compilation/correlation modeling with MT and temperature

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47 model data.

49 2. Bac-Man Geothermal System

50 The exploration of the Bac-Man geothermal field has been carried out since 1977. 51 Several studies have been carried out ranging from geology, geochemistry, geophysics 52 (gravity, magnetic, micro earthquake, magneto-telluric, temperature, etc.) with the aim of 53 determining the geothermal reservoir, caprocks, heat-source, the geothermal system model 54 and geothermal potential in the area. The total capacity of the power plant in the Bac-Man 55 area is 150 MWe generated from 4 generating units, namely: Unit I began operating in 1993 56 and unit II in 1995 which was named Palayang Bayan (2 x 55 Mwe), in 1996 the Cawayan 57 geothermal plant began operating (20 Mwe), and in 1998 the Unit III power plant in the 58 Botong area began operating (20 Mwe). The Bac-Man geothermal field is located in the 59 Podcol mountains 350 km southeast of Manila. The local fault system is known as the 60 Bac-Man Fault Zone (BFZ) (Dimabayao, Rowe and Barker, 2019). The fault zone is 61 indicated by a series of volcanoes that are mostly NW-SE (Figure 1). Structurally controlled 62 by a fault system which is believed to be an extension of the San Vicente Linao Fault (SVLF) 63 which is a stretch of the Philippine fault. The most prominent regional geological structure in 64 the area is the San Vicente-Linao Fault in the Northwest - Southeast (NW-SE) which slopes 65 across the Bico Peninsula (Resyes, Delfin and Bueza, 1995).



Figure 1. Bicol Peninsula Fault Area, San Vicente-Linao Fault (SVLF) (Lagmay, Tengonciang and Uy,
2005).

70 The Bac-Man geothermal system is divided into two regions, namely the western and 71 eastern parts of Bac-Man. The eastern part of Bac-Man is further divided into the North 72 Namito lowlands and the Podcol highlands, where the Podcol plateau can be divided into 73 eight geographic sectors, namely: Inang Maharang, Putting Bato, Palayan Bayan, Cawayan, 74 Tanawon, Osiao, Bangas and Botong (Resyes, Delfin and Bueza, 1995). Neutral chloride hot 75 springs with a temperature of 89 ° C - 96 ° C are found in the lowlands of Manito. Solfatara is 76 found in Cawayan and Pangas, while in West BacMan, cold to warm and cold SO4 springs 77 are found.

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The most recent volcanic event in the Pocdol highlands occurred more than 40 thousand years ago. This is related to the formation of the Tanawon and Cawayan craters, as well as the extrusion of the Botong and Pangas domes. The youngest volcanoes generally occur in subsurface areas with high temperatures, permeable formations, and active thermal manifestations (Figure 2) (Resyes, Delfin and Bueza, 1995).





Figure 2. Structure and manifestations at the Bac-Man geothermal field (Africa, 2013).

Layugan *et al.* (2005), analyzed and interpreted magneto-telluric data based on data from 1999 and 2001 which obtained the contour boundaries of the anomaly resistivity of the conductive zone as a reservoir area.

88 An iso-resistivity map at a depth of -1100 m from MSL which is correlated with the 89 geological structure in the area shows that the central part of the Bac-Man fault zone has a

- 90 low resistivity value (Tugawin *et al.*, 2015). The Bac-Man reservoir area covers an area of 26
- 91 36 km2, while the Kayabon reservoir located to the northwest of Bac-Man covers an area of



92 12 - 18 km2 (Figure 3).

Figure 3. An isoresistivity map at a depth of -1100 m from MSL which is correlated with the
geological structure in the area shows that the central part of the Bac-Man fault zone has a
low resistivity value (Tugawin *et al.*, 2015).

97 The results of this MT study support the results of previous MT research conducted by 98 Layugan *et al.* (2005), which found that the prospect area for the Bac-Man geothermal 99 reservoir is in Botong, Cawayan and Tikolob although the results show a larger area than the 100 results of previous MT studies.

101 Research related to reservoir temperature was carried out by Austria, 2008, in which the
102 eastern part of Palayan - Bayan has a high temperature of 326 ° C (Austria, 2008). The





Figure 4. Temperature contour map at a depth of -1000 meters from MSL (Austria, 2008)

108 **3.** Methods

109 The gravity data used in this study were 125 points which were the results of 110 measurements from 2009 to 2010 (Monasterial, 2015). Processing of gravity data includes: 111 determination of surface density using the Parasnis and Nettleton method, determination of 112 Bouguer anomaly, Bouguer anomaly filtering, Bouguer anomaly separation, horizontal and 113 vertical gradient analysis of Bouguer anomaly, and Bouguer anomaly modeling. Analysis and 114 interpretation were carried out to verify the Bac-Man geothermal reservoir area and structures 115 in the area. The analysis and interpretation are carried out by correlating the gravity model 116 with geological data, well data and other geophysical data. The research flow carried out is 117 shown in Figure 5.



119 **Figure 5**. Flowchart of research in determining the boundary of a geothermal reservoir

Density determination was carried out using the Parasnis method and the Nettleton method. The results of calculations using the Parasnis method get a value of 2.24 g/cc, while calculations using the Nettleton method for 2 (two) cross sections get a value of 2.35 g/cc and 2.38 g/cc. This density value when compared with geological conditions in the field has an appropriate value, so that in research for the calculation of Bouguer correction and modeling using a density of 2.35 g/cc.

126 To obtain Bouguer anomaly, gravity observation data is performed by gravity theoretic 127 correction at latitude φ , free air correction (free air correction), Bouguer correction and 128 terrain correction. The calculation of the gravity theorical correction at latitude φ uses the 129 International Gravity Formula 1980 equation (Wellenhof and Moritz, 2005):

130
$$g_{\varphi} = 978.032,7(1+5.3024\ 10^{-2}\ Sin^2\varphi - 5,8\ 10^{-6}\ Sin^2\ 2\phi \tag{1}$$

Meanwhile, for free air correction, the FAA equation = -0.308 h (mGall/m) is used which is obtained from the derivative of the Earth's normal gravity equation in the form of an ellipsoid, namely:

$$g_{\phi,h} = g_{\phi} + \frac{\partial g_{\phi}}{\partial h}h \tag{2}$$

135
$$\frac{\partial g_{\phi}}{\partial h} = -\frac{2g_{\phi}}{\alpha} (1 + f + m - 2f \sin^2 \phi) = -0.308 \, mGall/m \tag{3}$$

136 The Bouguer correction value is calculated using the approach model for the slab model,137 namely:

138
$$Bc = 2\pi G\rho h = 0.04193\rho h$$
 (4)

139 Where ρ is rock density (g/cc), h is height (m), and Bc is Bouguer correction (mGall).

- 140 Terrain corrections were calculated using a combination of the equations given by Nagy
- 141 (1966) and Kane (1962). In the calculation of terrain correction, the topographic model is
- 142 approached with a prism-shaped arrangement of objects measuring 1 km x 1 km with a height
- 143 in accordance with the topography of the area up to a radius of 50 km. The topographical data
- 144 used is DEM data taken from INA Geoportal.

146 4. Result and Discussions

The Bouguer anomaly in the Bac-Man geothermal field has a value of 18 - 41 mGall, with a high anomaly in the West to the Northwest and a little in the middle, while the anomaly is low in the middle. The high anomaly in the Northwest part is probably related to the heat-source of the geothermal system in the Bac-Man field, while the low anomaly in the middle part flanking the high anomaly is probably related to the presence of a reservoir in the area (Figure 6).



153

154 Figure 6. Map of the Bouguer anomaly of the Bac-Man geothermal field area

155 Spectral analysis was carried out to determine the boundary of the regional bouguer 156 anomaly and the residual area of the study. The results of this spectral analysis are then used

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157 to estimate the window width for gravity anomaly filtering. In general, a Fourier 158 transformation is to reconstruct/unravel an arbitrary wave into a sine wave with a variable 159 frequency where the sum of the sine waves is the original waveform (Bhattacharyya and Leu 160 (1977); Ghosh and Singh (2014)). Spectrum analysis is used to determine the depth of the 161 structure of the anomaly. In this study, the analysis of the spectrum analysis used Fourier 162 transform, where the results are used to estimate the width of the filtering window. The 163 spectrum analysis was carried out by making a cross section of the Bouguer anomaly as many 164 as 5 trajectories with a point interval of 250 m. The result of spectrum analysis from 5 lines 165 of Bouguer anomaly shows that the average regional Bouguer anomaly depth is 2500 meters. 166 Based on these results, the separation of regional Bouguer anomalies and residual Bouguer 167 anomalies is used a moving average filter with a window width of 5 km x 5 km.

168 The residual Bouguer anomaly map from the Bouguer anomaly filtering using the 169 moving average method with windows 5 km x 5 km is shown in Figure 7. Residual Bouguer 170 anomaly maps have values from -12 mGall to 9 mGall with high anomalies occupying the 171 eastern and central parts of the study area surrounded by low anomalies in the north, south 172 and east. This area which occupies a low anomaly is probably the geothermal reservoir area 173 of the Bac-Man field and this is in accordance with the magnetoteluric geophysical data as 174 well as well data. There are 3 low anomaly closures in the middle which are separated by 175 high anomalies. It can be interpreted that the geothermal reservoir in the Bac-Man field may 176 be divided into 3 reservoirs, namely the North, South and East. Reservoirs may be separated 177 from one another by fault structures or the presence of intrusion in the area.





179 Figure 7. Map of the Bouguer Residual anomaly of the Bac-Man geothermal field

To support the analysis of the Bouguer Residual anomaly in identifying: the existence of the fault structure, the lithological boundary and to generate the shallow effect anomaly, a vertical gradient analysis of the Bouguer Residual anomaly was carried out. Theoretically, this method is derived from Laplace's equation for surface gravity anomalies:

184
$$\nabla^2 \Delta g = \mathbf{0} \text{ atau } \frac{\partial^2 Ag}{\partial x^2} + \frac{\partial^2 Ag}{\partial y^2} + \frac{\partial^2 Ag}{\partial z^2} = 0 \tag{5}$$

$$\frac{\partial^2 \Delta g}{\partial z^2} = -\frac{\partial^2 \Delta g}{\partial x^2} + \frac{\partial^2 \Delta g}{\partial y^2}$$
(6)

In this study, the SVD anomaly gravity value is calculated using a filtering process,namely through convolution between anomaly gravity and a second vertical derivative filter.

$$\Delta Gsvd(\Delta x, \Delta y) = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \Delta g(x, y) F(x - \Delta x, v - \Delta y) dx dy$$
(7)

189 Where F is the second vertical derivative filter according to the above equation and ΔG is the 190 gravity anomaly.

191 The second vertical derivative filter used in this study is the Elkins (1951) type. The 192 SVD map compiled Bouguer Residual anomaly with low residual Bouguer anomaly contours 193 and the boundaries of the Bac-Man field geothermal prospects from the MT data is shown in 194 Figure 8.

195 Data compilation of low Bouguer Residual anomaly, Bouguer Residual anomaly SVD, 196 and reservoir prospect boundaries derived from magneto-telluric data were carried out to 197 identify the presence of a reservoir in the area. In general, geothermal reservoirs will have a 198 low Bouguer anomaly, because a good geothermal reservoir will have a high porosity value, 199 high permeability so that the reservoir rock will have a low density. The SVD map of 200 Bouguer Residual anomaly compiled with: Low Residual Bouguer anomaly, reservoir area 201 boundary derived from magneto-telluric data, and reservoir prospect area derived from 202 Bouguer Residual anomaly data is shown in Figure 8.



204 Figure 8. SVD map of the residual Bouguer anomaly overlaid with the reservoir area derived

from magnetotelluric data, and the reservoir prospect area derived from the Bouguer Residual anomaly data.

The figure shows that the reservoir locations of the four data provide the same results. The reservoir areas identified from the Bouguer Residual anomaly data and the Bouguer Residual SVD anomaly provide more detailed results, namely the Bac-Man geothermal reservoir is separated into 3 areas, namely: the Cawayan and Tanawon sectors in the South, the Boton sector which is located in the East, and the East sector. Palayang. This result is also supported by the high temperature in the area, where the highest temperature is in the Palayan sector.

The temperature distribution map at a depth of 1200 m from MSL shows that the center has a high temperature which is probably due to intrusion in the area. The existence of this intrusion is in accordance with the high residual gravity anomaly pattern in the area, where the intrusion is also a barrier from 3 reservoir locations in the area.

To obtain a model of the subsurface structure of the Bac-Man Geothermal Field in this study, a 3D inversion modeling of Bouguer Residual anomaly was carried out. The equation used and the calculation of 3D inversion modeling is the subsurface model approach which is composed of prisms with the amount according to: measurement area, data grid, thickness and depth of objects. The calculation of the gravity response for each prism block uses the Plouff (1976) equation:

$$g = G\Delta\rho \sum_{i=1}^{2} \sum_{j=1}^{2} \sum_{k=1}^{2} \mu_{ijk} \left[z_k \arctan \frac{x_i y_i}{z_k R_{ijk}} - x_i \log(R_{ijk} + y_i) - y_i \log(R_{ijk} + x_i) \right]$$
224

225 Where:
$$R_{ijk} = \sqrt{x_i^2 + y_j^2 + z_k^2}$$
; $\mu_{ijk} = (-1)^i (-1)^j (-1)^k$

The results of Bouguer Residual anomaly 3D inversion modeling using the Grav3D program are shown in Figure 9. The density distribution model of the 3D inversion modeling results shows the distribution of high and low density in the middle of the study area with a value of 2 g/cc to 2.9 g/cc. To obtain a subsurface structure model, the model is then carried out by slicing the selected paths which will be compared and correlated with models such as: temperature data models, magneto-telluric cross-sectional models, structural models and others.



233



The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the temperature cross-sectional model and the intrusion structure on the NW-SE trajectory is shown in Figure 10. The results of the 3D Bouguer anomaly residual inversion result which are correlated with the temperature section model

- 240 (Ramos, 2002) and the intrusion structure shows a good correlation, the presence of intrusion
- 241 is correlated with high density in the area. Likewise, the SVD section indicates that the
- intrusion is no single but separated into three parts according to the model derived from well
- 243 data and temperature data.



Figure 10. The cross section of the density distribution model resulting from the 3D Bouguer Residual anomaly inversion correlated with the temperature and intrusion structure cross-sectional model for the NW-SE trajectory.

The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the temperature cross-section model and the intrusion structure on the NE-SW trajectory is shown in Figure 11. The results of the 3D Bouguer anomaly residual inversion result which correlated the temperature cross-section model and

the intrusion structure show the existence of good correlation, the presence of intrusion correlates with high density in the area. Likewise, the SVD section indicates that the intrusion is no single but separated into two parts according to the model derived from well data and temperature data.



256

Figure 11. The cross section of the density distribution model resulted from the Bouguer Residual anomaly 3D inversion which is correlated with the temperature and intrusion structure cross-sectional model for the NE-SW trajectory.

The model is correlated with: the boundary of the geothermal reservoir area delineated from the MT results, the division of the geothermal sector, and the presence of volcanoes that control the geothermal system in the Bac-Man field. 263 The results of the correlation analysis between the cut density distribution models for 264 density 2.1 g/cc - 2.4 g/cc (interpreted as a geothermal reservoir) and the reservoir prospect 265 boundary derived from MT indicate that the reservoir locations have similarities, but the 266 reservoir area derived from the gravity model is more broad and in more detail. From the 267 gravity modeling shows that the geothermal reservoir in the Bac-Man field is divided into 3 268 areas, namely: the southern part (Cawayan and Tanawon sector), the eastern part (Boton 269 sector) and the northern part (Palayan -Inang Maharang sector) where each reservoir area is. 270 separated by the presence of fault structures or intrusions in the area. The division of the 3 271 reservoir areas is in accordance with the well data and the results of reservoir modeling 272 research in the area, where the geothermal reservoir area of the Bac-Man field is divided into 273 3 sectors, namely the south (Cawayan and Tanawon sectors), the east (Boton sector) and the 274 north (the North sector). Palayan-Inang Maharang) (Figure 12). The reservoir sectors are 275 separated by intrusion in the area, this is supported by the temperature of the reservoir in the 276 area which has a higher temperature than the surrounding area.



Figure 12. Density distribution model of Bouguer Residual anomaly 3D inversion modeling results shown for densities 2.1 g/cc - 2.4 g/cc. The model is correlated with: geothermal reservoir area boundaries delineated from the MT results, SVD = 0 maps, the division of geothermal areas, and the presence of volcanoes that control the geothermal system in the Bac-Man field.

285 5. Conclusions

Based on the results of gravity research on the Bac-Man geothermal field supported by geological data, well data, magneto-telluric data, and temperature data, several conclusions can be drawn regarding the Bac-Man geothermal field, namely:

- Bouguer anomaly in the study area has a value of 18 41 mGall, with a high anomaly in the West to the Northwest and a little in the middle, while the anomaly is low in the middle. From the analysis of the Bouguer anomaly spectrum, it is found that the Bouguer Regional and Residual anomaly limits are at a depth of 2500 m.
- 293
 2. Residual Bouguer anomaly calculated using a moving average filter with windows 5
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- 299 3. The density distribution model from the 3D inversion modeling shows that the 300 density distribution model cut for density 2.1 g/cc - 2.4 g/cc has a correlation with 301 the reservoir boundary derived from MT, but the reservoir area derived from the 302 gravity model is wider and more detailed. The results of gravity modeling show that 303 the geothermal reservoir in the Bac-Man field is divided into 3 areas, namely: the 304 southern part (Cawayan and Tanawon sector), the eastern part (Boton sector) and the 305 northern part (Palayan - Inang Maharang sector) where each reservoir area is. 306 separated by the presence of fault structures or intrusions in the area. The division of 307 the 3 reservoir areas is in accordance with the well data and the results of reservoir 308 modeling research in the area. The reservoir sectors are separated by intrusion in the

309		area. This is also supported by temperature data which shows that the area has a
310		higher temperature than the surrounding area.
311	4.	In order to obtain more detailed reservoir boundaries, especially the outer reservoir
312		boundaries, it is suggested to add more detailed gravity measurement points in the
313		North, East and South.
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322	Acknow	vledgements
323	We	e would like to thank all those who have helped in the implementation of this research,
324	especia	lly the Geophysical Engineering Department, Faculty of Engineering, University of
325	Lampu	ng.

327 **References**

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1	Reservoir Identification of Bac-Man Geothermal Field Based on
2	Gravity Anomaly Analysis and Modeling
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8	Abstract
9	The Bac-Man geothermal field is located between the provinces of Albay and Sorsogon
10	on the Bicol Peninsula. Several geophysical modeling methods have been carried out to
11	identify the Bac-Man field geothermal system. This research focuses on 3D modeling and
12	analysis of gravity data which aims to identify the geothermal reservoir in the Bac-Man field.
13	This research includes several things, such as spectrum analysis and the separation of
14	Bouger's anomalies, gradient analysis, and anomaly modeling. Based on the modeling results,
15	there are 3 low anomaly closures in the middle which are separated by high anomalies, which
16	can be interpreted that the geothermal reservoir in the Bac-Man field may be divided into 3
17	reservoirs, namely the southern part (Cawayan and Tanawon sector), the eastern part (Boton
18	sector) and the northern part (Palayan - Inang Maharang sector) where each reservoir area is
19	separated by the presence of a fault or intrusion structure.

21 Keywords: Bac-Man field, Gravity, Modeling, Reservoir, Geothermal

23 1. Introduction

The Bacon - Manito (Bac-Man) geothermal field is located in the Podcol mountains on the island of Luzon, about 350 km southeast of Manila. The exploration of the Bac-Man geothermal field has been carried out since 1977. The reservoir model has 23 km², 1500 m depth, and a temperature of 240°C to 320°C. Based on geological, geochemical, and geophysical data, a conceptual model of the Bac-Man geothermal system has been made [1].

Tugawin *et al.* (2015) [2] carried out 2D magneto-telluric (MT) inversion modeling in the Bac-Man field, which shows the existence of 3 (three) geothermal reservoir areas, namely Palayan Bayan, Tikolob, and Malobago areas. These results follow the results of data interpretation of resistivity vertical electric sounding (VES) and Schlumberger Resistivity Traversing (SRT) [2]. The reservoir in the Palayan Bayan area has been in production since 1993, while the Tikolob prospect is on the west side of the Bac-Man field, separate from the Palayan Bayan system [2].

Espartinez and See (2005) [3] have conducted geochemical research in the Bac-Man Geothermal field to determine changes in reservoirs from their chemical properties. It is done to maintain the continuity of steam production in the Bac-Man field considering that of the 3 geothermal reservoir sectors, the Palayan Bayan and Cawayan sectors are still producing while the Boton sector has not been producing since 2009 [4].

In this research, gravity data processing and modeling will be carried out to determine the geothermal reservoir in the Bac-Man field (Palayan Bayan, Cawayan, and Boton). The results of the subsurface model derived from anomaly gravity analysis and the density distribution model resulting from compilation/correlation modeling with MT and temperature model data.
47 2. Bac-Man Geothermal System

48 The exploration of the Bac-Man geothermal field has been carried out since 1977. 49 Several studies have been carried out, ranging from geology, geochemistry, geophysics 50 (gravity, magnetic, micro earthquake, MT, temperature, etc.) to determine the geothermal 51 reservoir, caprocks, heat-source, the geothermal system model, and geothermal potential in 52 the area. The total capacity of the power plant in the Bac-Man area is 150 MWe generated 53 from 4 generating units, namely: Unit I began operating in 1993 and unit II in 1995 which 54 was named Palayang Bayan (2 x 55 Mwe), in 1996 the Cawayan geothermal plant began 55 operating (20 Mwe), and in 1998 the Unit III power plant in the Botong area began operating 56 (20 Mwe). The Bac-Man geothermal field is located in the Podcol mountains 350 km 57 southeast of Manila. The local fault system is known as the Bac-Man Fault Zone (BFZ) [5]. 58 The fault zone is indicated by a series of volcanoes that are mostly NW-SE (Figure 1). 59 Structurally controlled by a fault system, which is believed to be an extension of the San 60 Vicente Linao Fault (SVLF), is a stretch of the Philippine fault. The most prominent regional 61 geological structure in the area is the San Vicente-Linao Fault in the Northwest - Southeast 62 (NW-SE), which slopes across the Bico Peninsula [6].



63

64

Figure 1. Bicol Peninsula fault area, San Vicente-Linao fault (SVLF) [7].

65

The Bac-Man geothermal system is divided into two regions, namely the western and eastern parts of Bac-Man. The eastern part of Bac-Man is further divided into the North Namito lowlands and the Podcol highlands. The Podcol plateau can be divided into eight geographic sectors: Inang Maharang, Putting Bato, Palayan Bayan, Cawayan, Tanawon, Osiao, Bangas, and Botong [6]. Neutral chloride hot springs with a temperature of 89 °C - 96 °C are found in the lowlands of Manito. In Cawayan and Pangas, Solfatara is found in West Bac-Man, cold to warm and cold SO₄ springs are found.

73 The most recent volcanic event in the Pocdol highlands occurred more than 40 thousand 74 years ago. That is related to the formation of the Tanawon and Cawayan craters and the Botong and Pangas domes extrusion. The youngest volcanoes generally occur in subsurface
areas with high temperatures, permeable formations, and active thermal manifestations
(Figure 2) [6].



79

Figure 2. Structure and manifestations at the Bac-Man geothermal field [8].

Layugan *et al.* (2005) [9] analyzed and interpreted MT data based on data from 1999 and 2001, which obtained the contour boundaries of the conductive zone's anomaly resistivity as a reservoir area.

An iso-resistivity map at a depth of -1100 m from MSL, which is correlated with the geological structure in the area, shows that the central part of the Bac-Man fault zone has a low resistivity value [2]. The Bac-Man reservoir area covers 26 - 36 km2, while the Kayabon reservoir located to the northwest of Bac-Man covers an area of 12 - 18 km2 (Figure 3).



87

Figure 3. An isoresistivity map at a depth of -1100 m from MSL which is correlated with the geological structure in the area, shows that the central part of the Bac-Man fault zone has a low resistivity value [2].

The results of this MT study support the results of previous MT research conducted by Layugan *et al.* (2005) [9], which found that the prospect area for the Bac-Man geothermal reservoir is in Botong, Cawayan, and Tikolob. However, the results show a larger area than the results of previous MT studies.

Research related to reservoir temperature was carried out by Austria (2008) [1], in which
the eastern part of Palayan - Bayan has a high temperature of 326 ° C. The temperature
contour map at a depth of -1000 meters from MSL is shown in Figure 4.



Figure 4. Temperature contour map at a depth of -1000 meters from MSL [1].

102 **3.** Methods

103 The gravity data used in this study were 125 points: the results of measurements from 104 2009 to 2010 [10]. Processing gravity data includes determination of surface density using 105 the Parasnis and Nettleton method, the determination of Bouguer anomaly, Bouguer anomaly 106 filtering, Bouguer anomaly separation, and horizontal and vertical gradient analysis Bouguer 107 anomaly, and Bouguer anomaly modeling. Analysis and interpretation were carried out to 108 verify the Bac-Man geothermal reservoir area and structures in the area. The analysis and 109 interpretation are carried out by correlating the gravity model with geological data, well data, 110 and other geophysical data. The research flow carried out is shown in Figure 5.





112 **Figure 5**. Flowchart of research in determining the boundary of a geothermal reservoir

113 Density determination was carried out using the Parasnis method and the Nettleton

method. The calculations using the Parasnis method get a value of 2.24 g/cc, while calculations using the Nettleton method for 2 (two) cross-sections get a value of 2.35 g/cc and 2.38 g/cc. When compared with geological conditions in the field, this density value has an appropriate value, so that in research for the calculation of Bouguer correction and modeling using a density of 2.35 g/cc.

Bouguer anomaly obtained, gravity observation data is performed by gravity theoretic correction at latitude φ , free air correction (free air correction), Bouguer correction and terrain correction. The calculation of the theoretical gravity correction at latitude φ uses the International Gravity Formula 1980 equation (equation 1) [11]:

123
$$g_{\varphi} = 978.032,7(1+5.3024\ 10^{-3}\ Sin^2\varphi - 5,8\ 10^{-6}\ Sin^2\ 2\phi \tag{1}$$

Meanwhile, for free air correction, the FAA equation = -0.308 h (mGal/m) is used, which is obtained from the derivative of the Earth's normal gravity equation in the form of an ellipsoid, namely:

$$g_{\phi,h} = g_{\phi} + \frac{\partial g_{\phi}}{\partial h}h \tag{2}$$

128
$$\frac{\partial g_{\phi}}{\partial h} = -\frac{2g_{\phi}}{a}(1+f+m-2f\sin^2\phi) = -0.308 \, mGall/m \tag{3}$$

129 The Bouguer correction value is calculated using the approaching model for the slab130 model, namely:

131
$$Bc = 2\pi G\rho h = 0.04193\rho h$$
 (4)

132 Where ρ is rock density (g/cc), h is the height (m), and Bc is Bouguer correction (mGal).

133 Terrain corrections were calculated using a combination of the equations given by Nagy

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- 134 (1966) [12] and Kane (1962) [13]. In the calculation of terrain correction, the topographic
- 135 model is approached with a prism-shaped arrangement of objects measuring 1 km x 1 km
- 136 with a height following the topography of the area up to a radius of 50 km. The topographical
- 137 data used is DEM data taken from INA Geoportal.

139 **Result and Discussions** 4.

140 The Bouguer anomaly in the Bac-Man geothermal field has a value of 18 - 41 mGal, 141 with a high anomaly in the west to the northwest and a little in the middle, while the anomaly 142 is low in the middle. The high anomaly in the Northwest part is probably related to the 143 geothermal system's heat source in the Bac-Man field. The low anomaly in the middle part 144 flanking the high anomaly is probably related to a reservoir in the area (Figure 6).





146

Figure 6. Map of the Bouguer anomaly of the Bac-Man geothermal field area

147 Spectral analysis was carried out to determine the boundary of the regional Bouguer 148 anomaly and the residual area of the study. The results of this spectral analysis are then used to estimate the window width for gravity anomaly filtering. In general, a Fourier 149

150 transformation is to reconstruct/unravel an arbitrary wave into a sine wave with a variable 151 frequency where the sum of the sine waves is the original waveform [14-15]. Spectrum 152 analysis is used to determine the depth of the structure of the anomaly. In this study, the 153 analysis of the spectrum analysis used Fourier transform, where the results are used to 154 estimate the width of the filtering window. The spectrum analysis was carried out by making 155 a cross-section of the Bouguer anomaly as many as 5 trajectories with a point interval of 250 156 m. The result of spectrum analysis from 5 lines of Bouguer anomaly shows that the average 157 regional Bouguer anomaly depth is 2500 meters. Based on these results, the separation of 158 regional Bouguer anomalies and residual Bouguer anomalies is a moving average filter with a 159 window width of 5 km x 5 km.

160 The residual Bouguer anomaly map from the Bouguer anomaly filtering using the 161 moving average method with windows 5 km x 5 km is shown in Figure 7. Residual Bouguer 162 anomaly maps have values from -12 mGal to 9 mGal, with high anomalies occupying the 163 eastern and central parts of the study area surrounded by low anomalies in the north, south, 164 and east. This area that occupies a low anomaly is probably the geothermal reservoir area of 165 the Bac-Man field. This is following the MT geophysical data and well data. There are 3 low 166 anomaly closures in the middle, which are separated by high anomalies. It can be interpreted 167 that the geothermal reservoir in the Bac-Man field may be divided into 3 reservoirs, namely 168 the north, south, and east. Reservoirs may be separated from one another by fault structures 169 or the presence of intrusion in the area.



171 **Figure 7**. Map of the Bouguer Residual anomaly of the Bac-Man geothermal field

To support the Bouguer Residual anomaly analysis to generate the existence of the fault structure, the lithological boundary, and the shallow effect anomaly, a vertical gradient analysis of the Bouguer Residual anomaly was carried out. Theoretically, this method is derived from Laplace's equation for surface gravity anomalies:

176
$$\nabla^2 \Delta g = 0 \quad \text{atau} \quad \frac{\partial^2 \Delta g}{\partial x^2} + \frac{\partial^2 \Delta g}{\partial y^2} + \frac{\partial^2 \Delta g}{\partial z^2} = 0 \tag{5}$$

$$\frac{\partial^2 \Delta g}{\partial z^2} = -\frac{\partial^2 \Delta g}{\partial x^2} + \frac{\partial^2 \Delta g}{\partial y^2}$$
(6)

In this study, the SVD anomaly gravity value is calculated using a filtering process,namely through convolution between anomaly gravity and a second vertical derivative filter.

180

$$\Delta Gsvd(\Delta x, \Delta y) = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \Delta g(x, y) F(x - \Delta x, v - \Delta y) dxdy$$
(7)

181 Where F is the second vertical derivative filter according to the above equation and ΔG is the 182 gravity anomaly.

This study's second vertical derivative filter is the Elkins (1951) [16] type. The SVD map compiled Bouguer Residual anomaly with low residual Bouguer anomaly contours. The boundaries of the Bac-Man field geothermal prospects from the MT data are shown in Figure 8.

187 Data compilation of low Bouguer Residual anomaly, Bouguer Residual anomaly SVD, 188 and reservoir prospect boundaries derived from MT data were carried out to identify a 189 reservoir in the area. In general, geothermal reservoirs will have a low Bouguer anomaly 190 because a good geothermal reservoir will have a high porosity value and high permeability. 191 The reservoir rock will have a low density. The SVD map of Bouguer Residual anomaly 192 compiled with: Low Residual Bouguer anomaly, reservoir area boundary derived from MT 193 data, and reservoir prospect area derived from Bouguer Residual anomaly data is shown in 194 Figure 8.





14

197 from MT data and the reservoir prospect area derived from the Bouguer Residual anomaly198 data.

The figure shows that the reservoir locations of the four data provide the same results. The reservoir areas identified from the Bouguer Residual anomaly data and the Bouguer Residual SVD anomaly provide more detailed results, namely the Bac-Man geothermal reservoir is separated into 3 areas, namely: the Cawayan and Tanawon sectors in the south, the Boton sector, which is located in the east, and the East sector Palayang. This result is also supported by the high temperature in the area, where the highest temperature is in the Palayan sector.

The temperature distribution map at a 1200 m from MSL shows that the center has a high temperature which is probably due to intrusion in the area. This intrusion follows the high residual gravity anomaly pattern in the area, where the intrusion is also a barrier from 3 reservoir locations in the area.

A 3D inversion modeling of Bouguer Residual anomaly was carried out to obtain a structural model of the Bac-Man Geothermal Field. The equation used and the calculation of 3D inversion modeling is the subsurface model approach composed of prisms with the amount according to measurement area, data grid, thickness, and depth of objects. The calculation of the gravity response for each prism block uses the Plouff (1976) equation [17]:

$$g = G\Delta\rho \sum_{i=1}^{2} \sum_{j=1}^{2} \sum_{k=1}^{2} \mu_{ijk} \left[z_k \arctan \frac{x_i y_i}{z_k R_{ijk}} - x_i \log (R_{ijk} + y_i) - y_i \log (R_{ijk} + x_i) \right]$$
(8)

216 Where:
$$R_{ijk} = \sqrt{x_i^2 + y_j^2 + z_k^2}$$
; $\mu_{ijk} = (-1)^i (-1)^j (-1)^k$

217 The results of Bouguer Residual anomaly 3D inversion modeling using the Grav3D

program are shown in Figure 9. The 3D inversion modeling results' density distribution model shows the distribution of high and low density in the middle of the study area with a value of 2 g/cc to 2.9 g/cc. The model is then carried out by slicing the selected paths, which will be compared and correlated with models such as temperature data models, MT cross-sectional models, structural models, and others to obtain a structural model.



223

Figure 9. The subsurface density distribution model from the 3D inversion modeling result of
Bouguer residual anomaly in Bac-Man geothermal field

The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the cross-sectional temperature model and the intrusion structure on the NW-SE trajectory, is shown in Figure 10. The 3D Bouguer anomaly residual inversion result correlated with the temperature section model [18], and the intrusion structure shows a good correlation. The presence of intrusion is correlated with high density in the area. Likewise, the SVD section indicates that the intrusion is no single but separated



into three parts according to the model derived from well data and temperature data.

233

Figure 10. The cross-section of the density distribution model resulting from the 3D Bouguer residual anomaly inversion correlated with the temperature and intrusion structure cross-sectional model for the NW-SE trajectory [18].

The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the temperature cross-section model and the intrusion structure on the NE-SW trajectory, is shown in Figure 11. The 3D Bouguer anomaly residual inversion result, which correlated the temperature cross-section model and the intrusion structure, shows a good correlation. The presence of intrusion correlates with high density in the area. Likewise, the SVD section indicates that the intrusion is no single but separated into two parts according to the model derived from well data and temperature data.



Figure 11. The cross-section of the density distribution model resulted from the Bouguer Residual anomaly 3D inversion which is correlated with the temperature and intrusion structure cross-sectional model for the NE-SW trajectory [18].

The model is correlated with: the boundary of the geothermal reservoir area delineated from the MT results, the division of the geothermal sector, and volcanoes that control the geothermal system in the Bac-Man field.

The results of the correlation analysis between the cut density distribution models for density 2.1 g/cc - 2.4 g/cc (interpreted as a geothermal reservoir) and the reservoir prospect boundary derived from MT indicate that the reservoir locations have similarities. However, the reservoir area derived from the gravity model is broader and in more detail. The gravity 255 modeling shows that the geothermal reservoir in the Bac-Man field is divided into 3 areas: 256 the southern part (Cawayan and Tanawon sector), the eastern part (Boton sector), and the 257 northern part (Palayan-Inang Maharang sector). Each reservoir area is separated by the 258 presence of fault structures or intrusions in the area. The division of the 3 reservoir areas 259 follows the well data and the results of reservoir modeling research in the area. That the 260 geothermal reservoir area of the Bac-Man field is divided into 3 sectors, namely the south 261 (Cawayan and Tanawon sectors), the east (Boton sector), and the north (the North sector). 262 Palayan-Inang Maharang) (Figure 12). The reservoir sectors are separated by intrusion in the 263 area. This is supported by the reservoir temperature in the area with a higher temperature than 264 the surrounding area.



266 Figure 12. The density distribution model of Bouguer Residual anomaly 3D inversion

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- 267 modeling results shown for densities 2.1 g/cc 2.4 g/cc. The model is correlated with:
- 268 geothermal reservoir area boundaries delineated from the MT results, SVD = 0 maps, the
- 269 division of geothermal areas, and volcanoes that control the geothermal system in the
- 270 Bac-Man field.
- 271

273 **5.** Conclusions

Based on the results of gravity research on the Bac-Man geothermal field supported by geological data, well data, MT data, and temperature data, several conclusions can be drawn regarding the Bac-Man geothermal field, namely:

- Bouguer anomaly in the study area has a value of 18 41 mGal, with a high anomaly
 in the west to the northwest and a little in the middle, while the anomaly is low in the
 middle. From the Bouguer anomaly spectrum analysis, it is found that the Bouguer
 Regional and Residual anomaly limits are at a depth of 2500 m.
- 281
 2. Residual Bouguer anomaly calculated using a moving average filter with windows 5
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- 287 3. The density distribution model from the 3D inversion modeling shows that the 288 density distribution model cut for density 2.1 g/cc - 2.4 g/cc correlates with MT 289 reservoir boundary. However, the reservoir area derived from the gravity model is 290 broader and more detailed. The results of gravity modeling show that the geothermal 291 reservoir in the Bac-Man field is divided into 3 areas: the southern part (Cawayan 292 and Tanawon sector), the eastern part (Boton sector), and the northern part (Palayan -293 Inang Maharang sector) where each reservoir area is separated by the presence of 294 fault structures or intrusions in the area. The division of the 3 reservoir areas follows 295 the well data and the results of reservoir modeling research in the area. The reservoir 296 sectors are separated by intrusion in the area. Temperature data also support this,

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297		showing that the area has a higher temperature than the surrounding area.
298	4.	Detailed reservoir boundaries, especially the outer reservoir boundaries, it is
299		suggested to add more detailed gravity measurement points in the north, east, and
300		south.
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1	Reservoir Identification of Bac-Man Geothermal Field Based on
2	Gravity Anomaly Analysis and Modeling
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8	Abstract
9	The Bac-Man gravity modeling was carried out to describe the geological structures that
10	control the geothermal system and estimate the area of the geothermal reservoir. A total of
11	125 data points were used to produce a complete Bouguer anomaly map of the area. The
12	gravity data are separated into regional and residual components to enhance the structural
13	features of the sedimentary and basement rocks in the study area. Gravity data were analyzed
14	using gradient interpretation techniques for edge detection, such as vertical descent. To
15	perform three-dimensional (3-D) modeling, a 5 \times 5 km volume and a depth of 2.5 km were
16	selected. This study presents the interpretation of various gravity anomaly maps and 3-D
17	inversion models. The interpretation of the vertical derivative of the gravity data indicates the
18	presence of a low gradient anomaly. The anomaly map is used to identify several faults or
19	intrusions compared to the faults or intrusions that are mapped. The 3-D model reveals that
20	there are 3 geothermal reservoirs, and the average block density value is 2.25 g/cc. These
21	reservoirs are spread over the Southern area (Cawayan and Tanawon Sector), East area
22	(Boton Sector), and Palayan-Inang Maharang area. Faults or rock intrusions are the limiting
23	factors for the three existing reservoirs. The results obtained from this study will lead to a
24	better understanding of the geothermal system in the study area, in particular, reservoir

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- 25 boundaries, and assist in future geothermal exploration.
- 26
- 27 Keywords: Bac-Man field, Gravity, Modeling, Reservoir, Geothermal
- 28

29 1. Introduction

30 Geothermal energy is a natural source of heat contained within the Earth, and it can be 31 extracted and used to generate electricity or for heating applications. A geothermal system 32 consists of three elements: a heat source, reservoir rock, and caprock [1]. The Bacon - Manito 33 (Bac-Man) geothermal field is located in the Podcol mountains on Luzon, about 350 km 34 southeast of Manila. The exploration of the Bac-Man geothermal field has been carried out 35 since 1977. The reservoir model has 23 km², 1500 m depth, and a temperature of 240°C to 36 320°C. Based on geological and geophysical data, a conceptual model of the Bac-Man 37 geothermal system has been made [2]. Tugawin et al. (2015) [3] carried out 2D 38 magneto-telluric (MT) inversion modeling in the Bac-Man field, which shows the existence 39 of 3 (three) geothermal reservoir areas, namely Palayan Bayan, Tikolob, and Malobago areas. 40 These results follow the results of data interpretation of resistivity vertical electric sounding 41 (VES) [4] and Schlumberger Resistivity Traversing (SRT). The reservoir in the Palayan 42 Bayan area has been in production since 1993, while the Tikolob prospect is on the west side 43 of the Bac-Man field, separate from the Palayan Bayan system [3].

44 The gravity method is applied to determine the geometry of the geothermal reservoir and 45 describe the geological structures that control the system. The gravity method is generally 46 used to describe the subsurface structures that control geothermal systems in geothermal 47 exploration. Certain geothermal reservoirs and their fluid content cause density differences 48 between the geothermal reservoir and the surrounding rock. The basis of the gravity method 49 is the density contrast in the rock. Gravity studies in various regions of the world have 50 yielded valuable results for geothermal exploration, such as the investigation of basement 51 topography in geothermal fields [5]; magma chambers and intrusion bodies associated with 52 heat source from the geothermal system; and faults delineation and fracture zones 53 corresponding to geothermal system reservoirs [6]. Therefore, the gravity method is one of 54 the most economical geophysical methods for modeling geothermal systems. The gravity 55 anomaly interpretation procedure consists of many techniques depending on the quality of the 56 data set and the purpose of the analysis [7]. Gravity anomaly maps are generally analyzed 57 using multiple linear transformations, directional derivative-based techniques, and inverse 58 modeling techniques. Gravity data in the form of a Bouguer anomaly map is used to describe 59 the study area's geological characteristics and subsurface structures. Observations of gravity 60 at the Earth's surface reflect the superimposed effects of broader and deeper mass variations 61 as well as shallower and more localized changes near the point of observation. This research 62 focuses on the 3D modeling of gravity data to estimate the geothermal reservoir extent and 63 understand the subsurface structure of the geothermal system.

65 2. Bac-Man Geothermal System

66 The exploration of the Bac-Man geothermal field has been carried out since 1977. 67 Several studies have been carried out, ranging from geology and geophysics (gravity, 68 magnetic, microearthquake, MT, temperature) to determine the geothermal reservoir, 69 caprocks, heat-source, the geothermal system model, and geothermal potential in the area. 70 The total capacity of the power plant in the Bac-Man area is 150 MWe generated from 4 71 generating units, namely: Unit I began operating in 1993 and unit II in 1995 which was 72 named Palayang Bayan (2 x 55 Mwe), in 1996 the Cawayan geothermal plant began 73 operating (20 Mwe), and in 1998 the Unit III power plant in the Botong area began operating 74 (20 Mwe). The Bac-Man geothermal field is located in the Podcol mountains 350 km 75 southeast of Manila. The local fault system is known as the Bac-Man Fault Zone (BFZ) [8]. 76 The fault zone is indicated by a series of volcanoes that are mostly NW-SE (Figure 1). 77 Structurally controlled by a fault system, which is believed to be an extension of the San 78 Vicente Linao Fault (SVLF), is a stretch of the Philippine fault. The most prominent regional 79 geological structure in the area is the San Vicente-Linao Fault in the Northwest - Southeast 80 (NW-SE), which slopes across the Bico Peninsula [9].



82 Figure 1. Map of the study area location and faults structure in the Bac-Man geothermal field

83

[3][10].

84

The Bac-Man geothermal system is divided into two regions, namely the western and eastern parts of Bac-Man. The eastern part of Bac-Man is further divided into the North Namito lowlands and the Podcol highlands. The Podcol plateau can be divided into eight geographic sectors: Inang Maharang, Putting Bato, Palayan Bayan, Cawayan, Tanawon, Osiao, Bangas, and Botong [9]. Neutral chloride hot springs with a temperature of 89 °C - 96 °C are found in the lowlands of Manito. In Cawayan and Pangas, Solfatara is found in West Bac-Man, cold to warm, and cold SO₄ springs are found.

The most recent volcanic event in the Pocdol highlands occurred more than 40 thousand years ago. That is related to the formation of the Tanawon and Cawayan craters and the Botong and Pangas domes extrusion. The youngest volcanoes generally occur in subsurface areas with high temperatures, permeable formations, and active thermal manifestations (Figure 2) [9]. 97

98



Figure 2. Structure and manifestations at the Bac-Man geothermal field [10].

Layugan et al. (2005) [11] analyzed and interpreted MT data based on 1999 and 2001,
which obtained the contour boundaries of the conductive zone's anomaly resistivity as a
reservoir area.

An iso-resistivity map at a depth of -1100 m from MSL, which is correlated with the geological structure in the area, shows that the central part of the Bac-Man fault zone has a low resistivity value [3]. The Bac-Man reservoir area covers 26 - 36 km2, while the Kayabon reservoir located to the northwest of Bac-Man covers an area of 12 - 18 km2 (Figure 3).



Figure 3. An isoresistivity map at a depth of -1100 m from MSL which is correlated with the geological structure in the area, shows that the central part of the Bac-Man fault zone has a low resistivity value [3].

110 The results of this MT study support the results of previous MT research conducted by 111 Layugan et al. (2005) [11], which found that the prospect area for the Bac-Man geothermal 112 reservoir is in Botong, Cawayan, and Tikolob. However, the results show a larger area than 113 the results of previous MT studies.

Research related to reservoir temperature was carried out by Austria (2008) [2], in which the eastern part of Palayan - Bayan has a high temperature of 326 ° C. The temperature contour map at a depth of -1000 meters from MSL is shown in Figure 4.

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Figure 4. Temperature contour map at a depth of -1000 meters from MSL [2].

119

121 **3.** Methods

122 The gravity data used in this study were 125 points: measurements from 2009 to 2010 123 [12]. Processing gravity data includes determination of surface density using the Parasnis and 124 Nettleton method, Bouguer anomaly, Bouguer anomaly filtering, Bouguer anomaly 125 separation, vertical gradient analysis Bouguer anomaly, and Bouguer anomaly modeling. 126 Analysis and interpretation were carried out to verify the Bac-Man geothermal reservoir area 127 and structures in the area. The analysis and interpretation are carried out by correlating the 128 gravity model with geological data, well data, and other geophysical data. The research flow 129 carried out is shown in Figure 5.



130

131 **Figure 5**. Flowchart of research in determining the boundary of a geothermal reservoir

132 Density determination was carried out using the Parasnis method and the Nettleton

method. The Parasnis method calculations get a value of 2.24 g/cc, while calculations using the Nettleton method for 2 (two) cross-sections get a value of 2.35 g/cc and 2.38 g/cc. When compared with geological conditions in the field, this density value has an appropriate value in research for calculating Bouguer correction and modeling using a density of 2.35 g/cc.

Bouguer anomaly obtained, gravity observation data is performed by gravity theoretic correction at latitude φ , free air correction (free air correction), Bouguer correction, and terrain correction. The calculation of the theoretical gravity correction at latitude φ uses the International Gravity Formula 1980 equation (equation 1) [13]:

141
$$g_{\varphi} = 978.032.7(1 + 5.3024 \, 10^{-3} \, Sin^2 \varphi - 5.8 \, 10^{-6} \, Sin^2 \, 2\phi \tag{1}$$

Meanwhile, for free air correction, the FAA equation = -0.308 h (mGal/m) is used, which is obtained from the derivative of the Earth's normal gravity equation in the form of an ellipsoid, namely:

$$g_{\phi,h} = g_{\phi} + \frac{\partial g_{\phi}}{\partial h}h \tag{2}$$

146
$$\frac{\partial g_{\varphi}}{\partial h} = -\frac{2g_{\varphi}}{a}(1+f+m-2f\sin^2\phi) = -0.308 \, mGall/m \tag{3}$$

147 The Bouguer correction value is calculated using the approaching model for the slab148 model, namely:

149 $Bc = 2\pi G\rho h = 0.04193\rho h$ (4)

150 Where ρ is rock density (g/cc), h is the height (m), and Bc is Bouguer correction (mGal).

151 Terrain corrections were calculated using a combination of the equations given by Nagy 152 (1966) [14] and Kane (1962) [15]. In the calculation of terrain correction, the topographic Submission Template to Journal of Applied Science and Engineering

model is approached with a prism-shaped arrangement of objects measuring 1 km x 1 km
with a height following the area's topography up to a radius of 50 km. The topographical data
used is DEM data taken from INA Geoportal [16].

156 The Bouguer anomaly contains contributions from regional trends resulting from the 157 presence of deep and large structures. The effect of this structure on the gravitational field 158 appears as a large wavelength anomaly, which masks the smaller and shallower effects. To 159 highlight the gravitational anomaly associated with the source of interest for this work, a 160 regional-remaining split of the Bouguer anomaly was performed using the open-source 161 Generic Mapping Tools (GMT) software to map and plot geographic data. This tool adjusts 162 the trend surface with the grid and calculates the residuals. The regional trend is removed by 163 using the trend-matched grid and removing the polynomial trend in the grid file.

164 Spectral analysis was carried out to determine the boundary of the regional Bouguer 165 anomaly and the residual area of the study. The results of this spectral analysis are then used 166 to estimate the window width for gravity anomaly filtering. In general, a Fourier 167 transformation is to reconstruct/unravel an arbitrary wave into a sine wave with a variable 168 frequency where the sum of the sine waves is the original waveform [17-18]. Spectrum 169 analysis is used to determine the depth of the structure of the anomaly. In this study, the 170 analysis of the spectrum analysis used Fourier transform [19], where the results are used to 171 estimate the width of the filtering window.

172
$$F(g) = 2\pi \gamma m \frac{e^{|k|(z_0 - z_1)}}{|k|}$$
(5)

173 The energy spectrum of the equation is:

174
$$E(k) = \frac{4\pi^2 \gamma^2 \rho^2}{|k|^2} e^{-2|k|z}$$
(6)

175
$$\log E(k) = \log(4\pi^2 \gamma^2 \rho^2) - 2|k|z - 2\log|k|$$
(7)
$$\log E(k) = \log A - 2|k|z \tag{8}$$

177 where Z_0 (depth point), Z_1 (depth mass) $Z_1 > Z_0$, $k = 2\pi/\lambda$ ((wave number), λ (wavelength), g

178 (gravity anomaly), ρ (density).

The separation of regional Bouguer anomalies and residual Bouguer anomalies is a moving average filter with a 5 km x 5 km window width. The calculation of the moving average is done by averaging the anomaly values for several points of gravity, as shown by the equation:

$$\Delta g_{Reg}(i,j) = \frac{(\Lambda g(i-n,j-n) + \dots + \Lambda g(i,j) + \dots + \Lambda g(i+n,j+n))}{N}$$
(9)

184 where $n = \frac{N-1}{2}$, and N must be an odd number.

185 This average is the regional anomaly, and the residual anomaly is obtained by 186 subtracting the data from the gravity measurement with the regional anomaly [20].

There are several methods for detecting the edges caused by fault structures or geological boundaries. Most of these methods are high-pass filters based on the horizontal and vertical derivatives of the gravity anomaly. One of these methods is the vertical gradient method, which has been used intensively to delineate contacts of density change from gravity data or pseudo gravity data. Theoretically, this method is derived from Laplace's equation for surface gravity anomalies [19][21]:

193
$$\nabla^2 \Delta g = 0 \text{ atau } \frac{\partial^2 \Delta y}{\partial x^2} + \frac{\partial^2 \Delta y}{\partial y^2} + \frac{\partial^2 \Delta y}{\partial z^2} = 0 \tag{10}$$

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$$\frac{\partial^2 \Delta g}{\partial z^2} = -\frac{\partial^2 \Delta g}{\partial x^2} + \frac{\partial^2 \Delta g}{\partial y^2}$$
(11)

194

In this study, the Second Vertical Derivative (SVD) anomaly gravity value is calculated
using a filtering process, namely through convolution between anomaly gravity and an SVD
filter [19].

198
$$\Delta Gsvd(\Delta x, \Delta y) = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \Delta g(x, y) F(x - \Delta x, v - \Delta y) dxdy$$
(12)

199 Where F is the SVD filter according to the above equation and ΔG is the gravity 200 anomaly.

Gravity inversion aims to determine the density distribution that would explain the measurements or the shape and dimensions of density variation. However, the inversion of the field gravity data is one of the most ambiguous problems in exploration geophysics studies [22]. Due to data noise and inhomogeneity of geological bodies, inversion of gravity data is usually fraught with difficulties even with reasonably accurate gravity measurements and data reduction.

A 3D inversion modeling of Bouguer Residual anomaly was carried out to obtain a structural model of the Bac-Man Geothermal Field with open source GRAV3D software [23][24]. The equation used and the calculation of 3D inversion modeling is the subsurface model approach composed of prisms with the amount according to measurement area, data grid, thickness, and depth of objects. The calculation of the gravity response for each prism block uses the Plouff (1976) equation [25]:

$$g = G\Delta\rho \sum_{i=1}^{2} \sum_{j=1}^{2} \sum_{k=1}^{2} \mu_{ijk} \left[z_k \arctan \frac{x_i y_i}{z_k R_{ijk}} - x_i \log \left(R_{ijk} + y_i \right) - y_i \log \left(R_{ijk} + x_i \right) \right]$$
(8)

214 Where:
$$R_{ijk} = \sqrt{x_i^2 + y_j^2 + z_k^2}$$
, $\mu_{ijk} = (-1)^i (-1)^j (-1)^k$

217 **Result and Discussions** 4.

218 The Bouguer anomaly in the Bac-Man geothermal field has a value of 18 - 41 mGal, 219 with a high anomaly in the west to the northwest and a little in the middle, while the anomaly 220 is low in the middle. The high anomaly in the Northwest part is probably related to the 221 geothermal system's heat source in the Bac-Man field. The low anomaly in the middle part 222 flanking the high anomaly is probably related to a reservoir in the area (Figure 6).





224

Figure 6. Map of the Bouguer anomaly of the Bac-Man geothermal field area

225 The spectrum analysis was carried out by cross-sectioning the Bouguer anomaly as 226 many as 5 trajectories with a point interval of 250 m (Figure 7). The spectrum analysis result 227 from 5 lines of Bouguer anomaly shows that the average regional Bouguer anomaly depth is

- 228 2500 meters (Figure 8). Based on these results, the separation of regional Bouguer anomalies
- and residual Bouguer anomalies is a moving average filter with a 5 km x 5 km window width.



Figure 7. Bouguer anomaly spectrum analysis to determine the depth limits of regional and

232 residual anomalies. The spectrum analysis results for line 2 get the depth limit value of the

residual regional Bouguer anomaly of 2483.9 = 2500 m.

230

234



Figure 8. Bouguer anomaly spectrum analysis to determine the depth limits of regional and residual anomalies. The spectrum analysis results for line 2 get the depth limit value of the residual regional Bouguer anomaly of 2504 = 2500 m.

The residual Bouguer anomaly map from the Bouguer anomaly filtering using the moving average method with windows 5 km x 5 km is shown in Figure 9. Residual Bouguer anomaly maps have values from -12 mGal to 9 mGal. High anomalies occupy the eastern and central parts of the study area, surrounded by low anomalies in the north, south, and east. This area that occupies a low anomaly is probably the geothermal reservoir area of the Bac-Man field. This is following the MT geophysical data and well data. There are 3 low anomaly closures in the middle, which are separated by high anomalies. It can be interpreted that the geothermal reservoir in the Bac-Man field may be divided into 3 reservoirs, namely the north, south, and east. Reservoirs may be separated from one another by fault structures or the presence of intrusion in the area.



Figure 9. Map of the Bouguer Residual anomaly of the Bac-Man geothermal field

This study's SVD filter is the Elkins (1951) [21] type. The SVD map compiled Bouguer Residual anomaly with low residual Bouguer anomaly contours. The boundaries of the Bac-Man field geothermal prospects from the MT data are shown in Figure 10.

253 Data compilation of low Bouguer Residual anomaly, Bouguer Residual anomaly SVD,

and reservoir prospect boundaries derived from MT data were carried out to identify a reservoir in the area. In general, geothermal reservoirs will have a low Bouguer anomaly because a good geothermal reservoir will have a high porosity value and high permeability. The reservoir rock will have a low density. The SVD map of Bouguer Residual anomaly compiled with: Low Residual Bouguer anomaly, reservoir area boundary derived from MT data, and reservoir prospect area derived from Bouguer Residual anomaly data is shown in Figure 10.



Figure 10. SVD map of the residual Bouguer anomaly overlaid with the reservoir area derived from MT data and the reservoir prospect area derived from the Bouguer Residual anomaly data.

The figure shows that the reservoir locations of the four data provide the same results. The reservoir areas identified from the Bouguer Residual anomaly data and the Bouguer Residual SVD anomaly provide more detailed results, namely the Bac-Man geothermal reservoir is separated into 3 areas, namely: the Cawayan and Tanawon sectors in the south, the Boton sector, which is located in the east, and the East sector Palayang. This result is also supported by the high temperature in the area, where the highest temperature is in the Palayan sector.

272	The temperature distribution map at 1200 m from MSL shows that the center has a high
273	temperature, probably due to intrusion. This intrusion follows the high residual gravity
274	anomaly pattern in the area, where the intrusion is also a barrier from 3 reservoir locations in
275	the area. Based on drilling data in the Bac-Man geothermal area, dike and intrusive rocks such
276	as monzogabbro, monzodiorite, and epidote have high-density characteristics (gabbro = 3.03
277	g/cc; amphibolite = 2.96 g/cc) [26]. These dike and intrusion zones have high temperatures
278	interpreted as heat sources of the Bac-Man geothermal system [27]. The high gravity anomaly
279	is associated with the presence of high density below the surface.

The results of Bouguer Residual anomaly 3D inversion modeling using the Grav3D program are shown in Figure 11. The 3D inversion modeling result, density distribution model, shows high and low density in the middle of the study area with a value of 2 g/cc to 2.9 g/cc. The model is then carried out by slicing the selected paths, which will be compared and correlated with temperature data models, MT cross-sectional models, structural models, and others to obtain a structural model.



Figure 11. The subsurface density distribution model from the 3D inversion modeling resultof Bouguer residual anomaly in Bac-Man geothermal field

The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the cross-sectional temperature model and the intrusion structure on the NW-SE trajectory is shown in Figure 12. The 3D Bouguer anomaly residual inversion result correlated with the temperature section model [27], and the intrusion structure shows a good correlation. The presence of intrusion is correlated with high density in the area. Likewise, the SVD section indicates that the intrusion is no single but separated into three parts according to the model derived from well data and temperature data.



Figure 12. The cross-section of the density distribution model resulting from the 3D Bouguer

residual anomaly inversion correlated with the temperature and intrusion structure cross-sectional model for the NW-SE trajectory [27].

The density distribution model resulting from the 3D Bouguer anomaly residual inversion which is correlated with the temperature cross-section model and the intrusion structure on the NE-SW trajectory is shown in Figure 13. The 3D Bouguer anomaly residual inversion result, which correlated the temperature cross-section model and the intrusion structure, shows a good correlation. The presence of intrusion correlates with high density in the area. Likewise, the SVD section indicates that the intrusion is no single but separated into two parts according to the model derived from well data and temperature data.

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Figure 13. The cross-section of the density distribution model resulted from the Bouguer Residual anomaly 3D inversion which is correlated with the temperature and intrusion structure cross-sectional model for the NE-SW trajectory [27].

The model is correlated with: the boundary of the geothermal reservoir area delineated from the MT results, the division of the geothermal sector, and volcanoes that control the geothermal system in the Bac-Man field.

The results of the correlation analysis between the cut density distribution models for density 2.1 g/cc - 2.4 g/cc (interpreted as a geothermal reservoir) and the reservoir prospect boundary derived from MT indicate that the reservoir locations have similarities. However, the reservoir area derived from the gravity model is broader and in more detail. The gravity

318 modeling shows that the geothermal reservoir in the Bac-Man field is divided into 3 areas: 319 the southern part (Cawayan and Tanawon sector), the eastern part (Boton sector), and the 320 northern part (Palayan-Inang Maharang sector). Each reservoir area is separated by the 321 presence of fault structures or intrusions in the area. The division of the 3 reservoir areas 322 follows the well data and the results of reservoir modeling research in the area. That the 323 geothermal reservoir area of the Bac-Man field is divided into 3 sectors, namely the south 324 (Cawayan and Tanawon sectors), the east (Boton sector), and the north (the North sector). 325 Palayan-Inang Maharang) (Figure 14). The reservoir sectors are separated by intrusion in the 326 study area and supported by the reservoir temperature in the area with a higher temperature 327 than the surrounding area.





- 330 modeling shows densities 2.1 g/cc 2.4 g/cc. The model is correlated with: geothermal
- 331 reservoir area boundaries delineated from the MT results, SVD = 0 maps, the division of
- 332 geothermal areas, and volcanoes that control the geothermal system in the Bac-Man field.

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335 5. Conclusions

Based on the results of gravity research on the Bac-Man geothermal field supported by geological data, well data, MT data, and temperature data, several conclusions can be drawn regarding the Bac-Man geothermal field, namely:

- Bouguer anomaly in the study area has a value of 18 41 mGal, with a high anomaly
 in the west to the northwest and a little in the middle, while the anomaly is low in the
 middle. From the Bouguer anomaly spectrum analysis, it is found that the Bouguer
 Regional and Residual anomaly limits are at a depth of 2500 m.
- Residual Bouguer anomaly calculated using a moving average filter with windows 5
 km x 5 km shows that the residual Bouguer anomaly has a value of -12 mGal to 9
 mGal with high anomalies occupying the eastern and central parts. There are 3 low
 anomaly closures in the middle that are separated by high anomalies. The
 geothermal reservoir in the Bac-Man field may be divided into 3 reservoirs: the
 North, South, and East.
- 349 3. The density distribution model from the 3D inversion modeling shows that the 350 density distribution model cut for density 2.1 g/cc - 2.4 g/cc correlates with the MT 351 reservoir boundary. However, the reservoir area derived from the gravity model is 352 broader and more detailed. The results of gravity modeling show that the geothermal 353 reservoir in the Bac-Man field is divided into 3 areas: the southern part (Cawayan 354 and Tanawon sector), the eastern part (Boton sector), and the northern part (Palayan -355 Inang Maharang sector) where each reservoir area is separated by the presence of 356 fault structures or intrusions in the area. The division of the 3 reservoir areas follows the well data and the results of reservoir modeling research in the area. The reservoir 357 358 sectors are separated by intrusion in the area. Temperature data also support this,

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359	showing that the area has a higher temperature than the surrounding area.
360	4. Detailed reservoir boundaries, especially the outer reservoir boundaries, suggest
361	adding more detailed gravity measurement points in the north, east, and south.
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