



Wai Selabung geothermal reservoir analysis based on gravity method

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ABSTRACT

Research has been conducted using the gravity method in the Wai Selabung area, South Ogan Kemiring Ulu Regency, South Sumatra Province, correlated with geological data, magnetotellurics, and geochemical data. This research aims to get structural patterns, subsurface models and identify the heat source and reservoir areas of the Wai Selabung geothermal system. This study uses the gravity method to model the subsurface, which is correlated with magnetotelluric and geochemical data to identify reservoir prospect areas. The results obtained from this research include residual anomalies in the research area showing the presence of a northwest-southeast trending fault structure by the main fault structure of this area trending northwest-southeast and slightly southwest-northeast. Analysis of the Second Vertical Derivative value of zero indicates the boundaries of the geothermal reservoir in the middle of the research area. The results of the 3D inversion modeling of the research area show that low density (2 to 2.15 g/cm³) indicates the location of the reservoir, medium-density values (2.2 to 2.4 g/cm³) are tertiary sandstone sedimentary. The high-density distribution value (2.5 to 2.9 g/cm³) indicates a potential heat source. And based on the analysis of the gravity method correlated with geological data, magnetotelluric, and geochemical data, the prospect area for the Wai Selabung geothermal reservoir, is around Teluk Agung, Perekkan, and Talang Tebat.

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INTRODUCTION

Indonesia has the potential for renewable energy, namely the most significant geothermal energy in the world. Indonesia's geothermal potential is 29,038 MW, and only 1,226 MW has been utilized or 4.2%. The National Energy Policy has targeted geothermal to support 5% of the national energy mix by 2025. However, geothermal has only contributed approximately 4% with slow development, so an investigation is needed to develop geothermal fields in Indonesia.

The Wai Selabung area is located on the Sumatran fault zone, where many high-temperature geothermal locations are found. Manifestations were found at the site in the form of hot springs on the banks of the Wai

Selabung River, in a tributary of the Wai Selabung river, and alterations that appeared near Suban hot springs (banks of the Wai Selabung river) (Kusnadi et al., 2011).

Several studies have been conducted in this area, whether related to geothermal activity or not, including Nasution et al. (1999), who made regional mapping using the Baturaja sheet system. Research related to geothermal activity has been carried out by the Center for Geological Resources from 2011 through an integrated survey of the Wai Selabung Geothermal Area (Kusnadi et al., 2011). The last is the magneto-telluric (MT) Survey conducted in 2020 (Ilmi et al., 2020). The combined survey results show a potential geothermal resource of 68 MWe in this area, and the prospect area is in the

southwest. Previous research revealed that the prospect area is only in the southwest, so further research is needed to identify the Wai Selabung geothermal area's structural pattern, heat source, and Wai Selabung geothermal reservoir (Hadi et al., 2011). Gravity surveys were conducted for 210 stations.

The gravity method is used to identify and describe subsurface geology based on variations in the earth's gravity field caused by differences in density between rocks. In the case of geothermal, the difference in rock density is a reference for gravity research. The heat source area and its accumulation below the earth's surface can cause density differences with the surrounding rock mass. Geological structures can form permeable zones that can escape to the surface so that manifestations appear on the surface. This study aims at a complete 3D inversion of the Bouguer anomaly, gradient analysis for structural analysis, and to determine the geochemical characteristics of the geothermal Wai Selabung area. Bouguer anomaly inversion modeling correlated with MT, geological, and geochemical data.

Geology

The structure of the study area is included in the volcanic environment of the magmatic pathway in southern Sumatra. The formation of the geothermal system in the Wai Selabung region is influenced by volcanic and tectonic activity in the direction of the Sumatran fault. In general, the rocks that make up the study area consist of volcanic rocks in the southwest and clastic rocks in the northeast of the Tertiary age (Hadi et al., 2011).

The geological structure in the Wai Selabung area is dominated by a northwest-southeast direction cut by a fault in a southwest-northeast and a north-south. In general, the research area consists of volcanic rocks (Figure 1). Based on the physical characteristics and age of the rocks, the rock units are obtained in order from old to young; Akar Jangkang Lava, Sandstone, Asadimana Lava, Pematang Gong Lava, Old Breccia, Ranau Pyroclastic Flow, Laai Lava, Bengkok Lava, Pandan Lava, Gedang Lava, Peraan Lava, Tebat Gayat Lava, and Alluvium Deposits (Kusnadi et al., 2011).

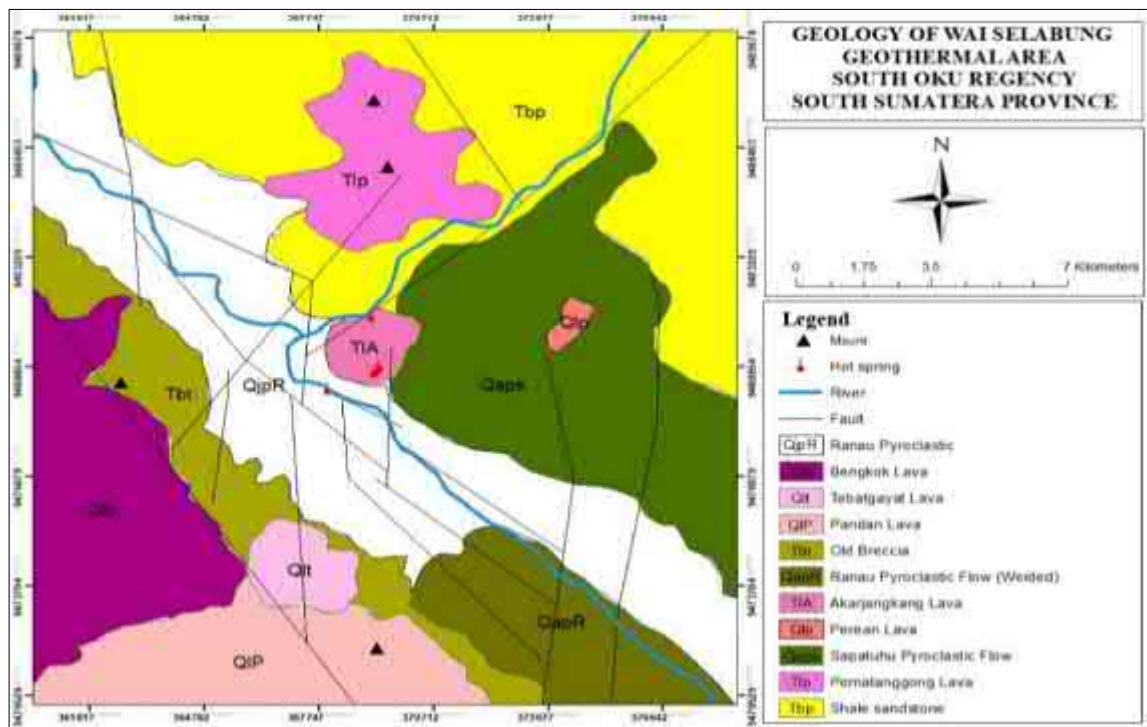


Figure 1. Geological map of the Wai Selabung geothermal area (modified from PSDG, 2011)

Surface Manifest

Based on the results of an integrated investigation in 2011, symptoms of geothermal appearances are characterized by the arrival of hot springs, which are divided into three groups of geothermal manifestations (Figure 2), namely

manifestations on the banks of the Wai Selabung River (2 groups), in Wai Selabung tributaries (2 groups). Alterations that occur on the banks of the Wai Selabung River appear near Suban hot springs (banks of the Wai Selabung river) (Kusnadi et al., 2011).

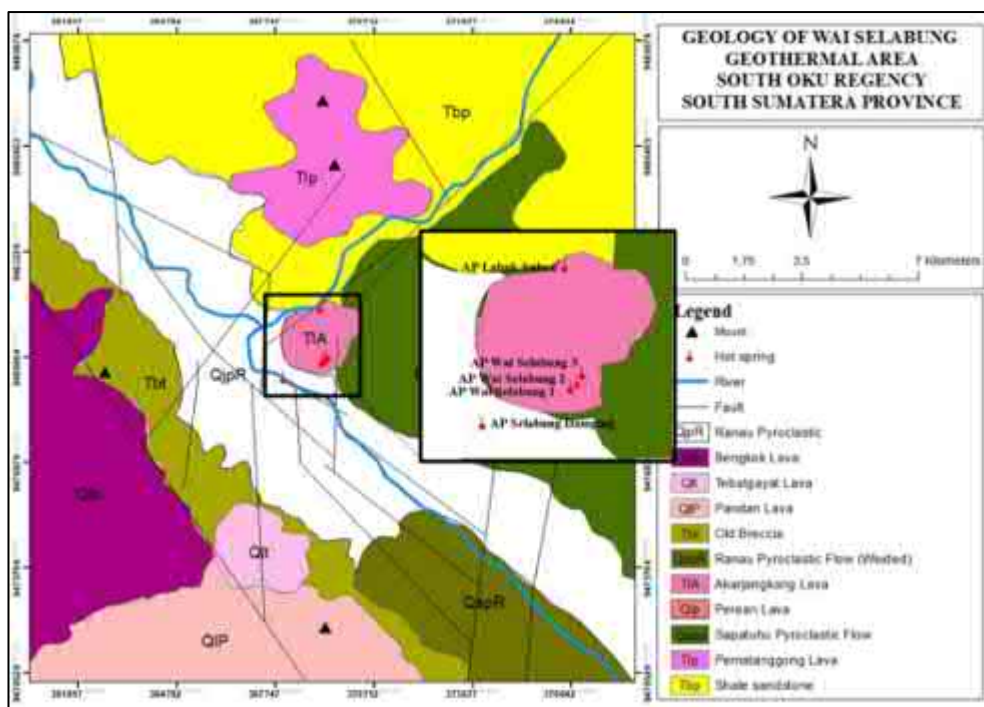


Figure 2. Wai Selabung Regional Manifestation Map (modification from PSDG, 2011)

The physical characteristics of the hot springs are shown in Table 1. From the pH data and the temperature of the manifestations, it can be seen that the highest temperature of the hot springs is in Wai Selabung Hot Water I, which is 92.5°C, the lowest is in Wai Selabung III Hot Water, which is 40.2°C. Meanwhile, the highest pH is found in Wai Selabung II Hot Water,

which is 9.43, and the lowest pH is in Selabung Damping Hot Water, 8.19. The degree of acidity (pH) is generally alkaline (8.19 – 9.47). Strong electrolytes have strong electrical conductivity because they contain a more significant number of ions than weak electrolytes. In Table 1, it can be seen that the smaller the pH value will be, the greater the power of the electricity.

Table 1. Geothermal Manifestations of the Wai Selabung Area.

No	Manifestation	Manifestation temperature (°C)	Air temperature (°C)	pH	DHL (μS/cm)
1	Wai Selabung 1 hot springs	92,5	25,43	9,43	2700
2	Wai Selabung II hot springs	89,3	29,9	9,47	2130
3	Wai Selabung III hot springs	40,2	40,2	8,38	686
4	Lubuk Suban hot springs	68,1	27	8,92	1196
5	Selabung Damping hot springs	44,4	25,1	8,19	715

Based on previous research, it was shown that the prospect area is only located in the southwest. Further research is needed to identify the Wai Selabung geothermal area's structural pattern, heat source, and reservoir Wai Selabung geothermal area using 3D inversion of the gravity anomaly. Furthermore, the gravity, geology, MT, and geochemical data compilation also supports identifying geothermal reservoir prospect areas.

METHODS

The gravity method was used to detect geothermal areas. The gravity method is a geophysical method based on the acceleration of gravity on the earth's surface. Changes influence variations in gravity values in the lateral density of the subsurface rock around the measurement point. This method identifies and describes subsurface geological structures based on variations in the earth's gravity field caused by differences in density between rocks.

In this study, a complete 3D inversion of the Bouguer anomaly was carried out, gradient analysis for structural analysis, and to determine the geochemical characteristics of the geothermal area of Wai Selabung. Bouguer anomaly inversion modeling correlated with MT, geological, and geochemical data.

The gravity data used in this study is secondary data from measurements carried out by the Center for Geological Resources in 2011 in the form of Complete Bouguer Anomaly. Gravity data was conducted gridding data and contouring to produce the Bouguer Anomaly Map using Surfer 16 software. Spectrum analysis was carried out in several paths to determine the depth of regional and residual anomalies, so it is known that the window width is used for the anomaly Bouguer filter. Spectrum analysis

was performed by Fourier transform (FFT) of trajectories representing the research area in a specific spatial domain (Δx) (Blakely, 1996).

$$F\left(\frac{1}{R}\right) = 2\pi \frac{e^{i|k|(z_0-z_1)}}{|k|} \quad (1)$$

Note $z_1 > z_0, |k| \neq 0$

The separation of the Bouguer anomaly is done with a moving average filter. Moving average is done by averaging the anomaly value. This average is the regional anomaly, while the residual anomaly is obtained by subtracting the data from the gravity measurement with the regional anomaly.

$$\Delta g_{Reg}(i, j) = \frac{(\Delta g(i-n, j-n) + \dots + \Delta g(i, j) + \dots + \Delta g(i+n, j+n))}{N} \quad (2)$$

Note $n = \frac{N-1}{2}$, and N must be an odd number.

Second Vertical Derivative (SVD) analysis of the Bouguer anomaly was carried out to obtain fault structures, lithological boundaries, and sources of anomalies originating close to the surface. The zero contours flanked by high and low anomaly values indicate the presence of a fault structure in the area.

$$\frac{\partial^2 \Delta g}{\partial z^2} = -\left(\frac{\partial^2 \Delta g}{\partial y^2} + \frac{\partial^2 \Delta g}{\partial x^2}\right) \quad (3)$$

3D inversion modeling of residual Bouguer anomaly was carried out to obtain a subsurface density distribution model. 3D inversion modeling using Grablox 1.7 software.

Based on the analysis of gravity, MT, geochemistry, and geology of the research area, a compilation map is formed showing the reservoir prospect area of the study area. The research data processing process can be seen in the following flow chart.

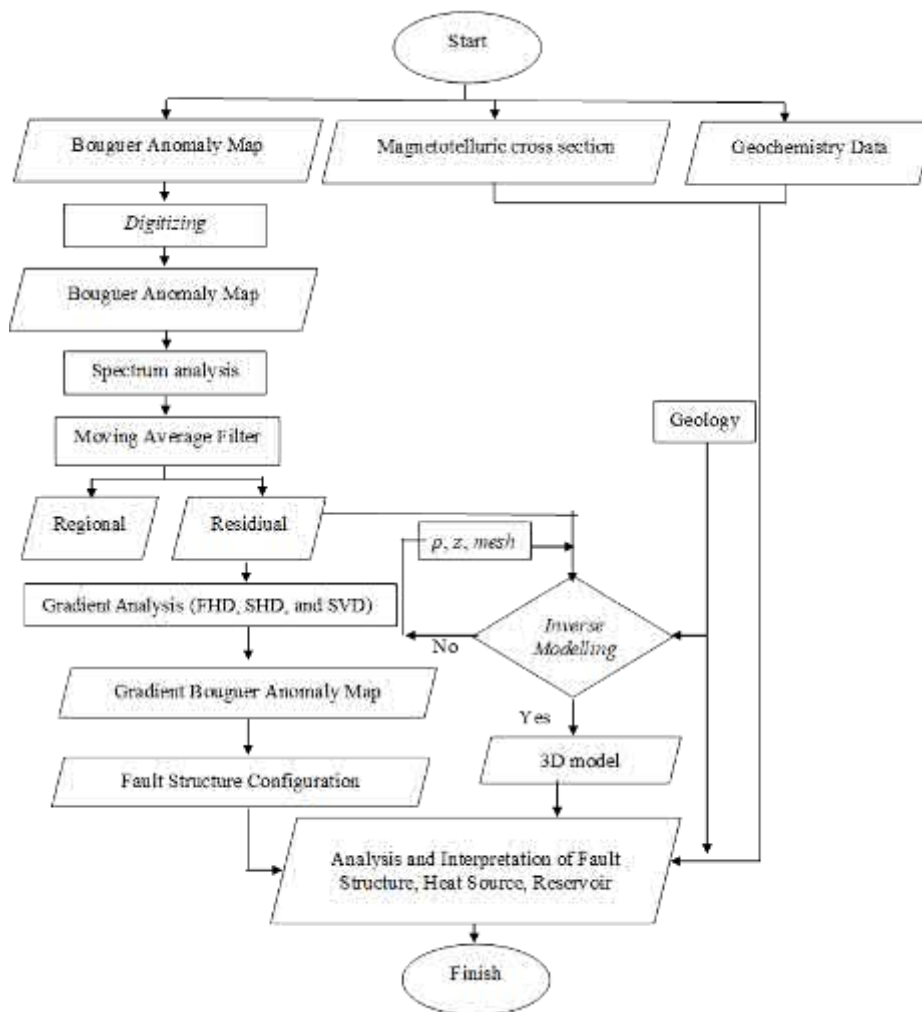


Figure 3. Flowchart of the Wai Selabung Geothermal Reservoir

RESULTS AND DISCUSSION

Bouguer Anomaly

The Bouguer anomaly value of the Wai Selabung area is the result of data processing of the Geological Resources research group from field measurements with several long trajectories from northwest to southeast, namely from Sinamarga to Talang Najam and from southwest to the northeast, namely the Wai Selabung area which is then processed to obtain Bouguer anomaly map.

Based on the research of Wahyu et al. (2019), the estimation of rock density was carried out using qualitative analysis and quantitative analysis. Qualitative analysis was carried out by measuring the density of rock samples from the investigation site at the Geological Resources Center laboratory. In contrast, the quantitative analysis used was

the Parasnis method. The laboratory analysis results show that the average value of the rock in the Wai Selabung geothermal research area is around 2.60 grams/cm³ (Wahyu et al., 2019).

The Bouguer anomaly map and the distribution of gravity measurement points in Figure 4 show that the study area has varying anomaly values ranging from 4 to 46 mGal. The high Bouguer anomaly value ranges from 36 mGal to 46 mGal, which is in the northwest, north, and northeast of the study area. In comparison, the low Bouguer anomaly value ranges from 4 mGal to 22 mGal, which is in the middle of the study area flanked by moderate gravity values with contour values of 22 mGal to 36 mGal in the Sinamarga and Kotadalam areas.

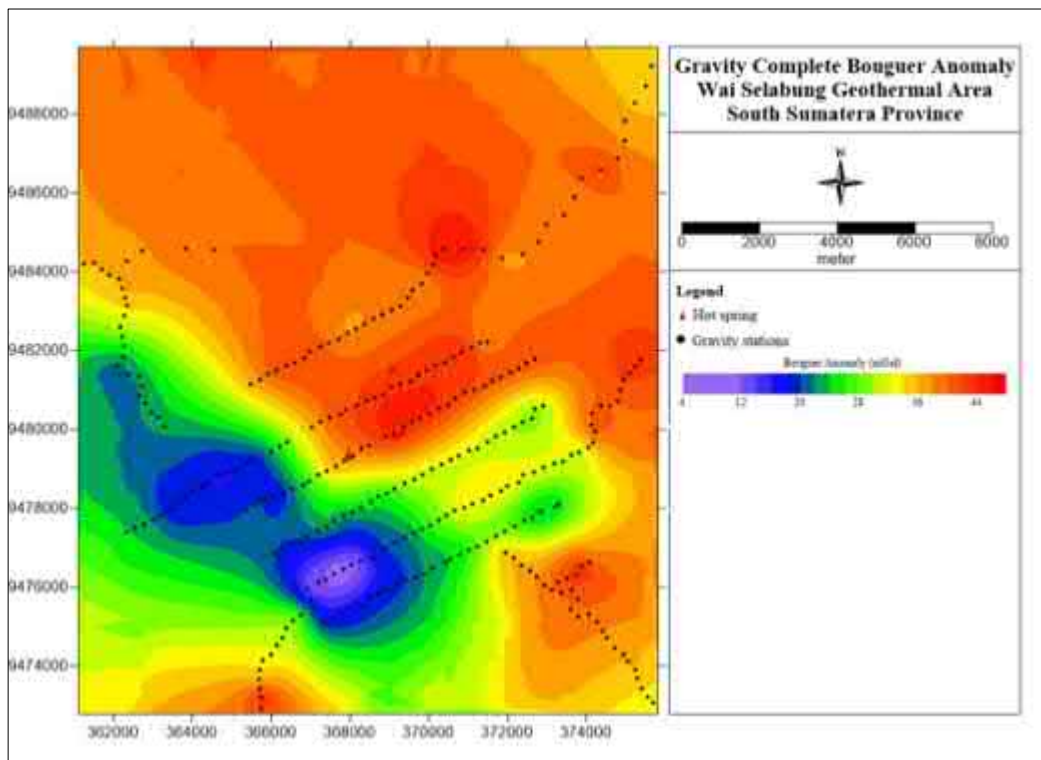


Figure 4. Bouguer Anomaly Map of the Study Area

Residual Anomaly

Residual anomaly is obtained from reducing the total anomaly value with the regional anomaly of the research area to bring up the residual anomaly value. The residual anomaly map shows a more complex pattern than the regional map and the Bouguer anomaly map. The residual anomaly map depicts an anomaly with a shorter wavelength that reflects the effect of a shallower object. The residual anomaly map of the study area has a residual anomaly value between -10 to 8 mGal. The residual anomaly map can be seen in Figure 5.

Low negative residual anomaly values range from -10 mGal to -4 mGal forming a closure in Kota Dalam, Teluk Agung, and Pematang Gong. The low anomaly is located in the center of the study area. Medium residual gravity values ranging from -3 mGal to 3 mGal dominate the investigation area in the north, south, east, and west directions.

Based on the research of Wahyu et al. (2019), the high anomaly was very clearly

demarcating the low anomaly zone so that the high anomaly zone, which was near from the manifestation. This shows that deep structures probably cause this residual anomaly with high complexity (Wahyu et al., 2019).

High residual gravity values with values >4 mGal poles in the middle of the study area between Pematang Gong and Teluk Agung where geothermal manifestations are found and in the Pematang Petang area to the southeast of the study area.

The residual anomaly shows a contour lination pattern in a northwest-southeast direction and shows positive and negative anomalous closures in the center of the study area. Where there are manifestations, indicates the presence of a northwest-southeast trending fault structure by the main fault structure of this area trending northwest-southeast and slightly southwest-northeast.

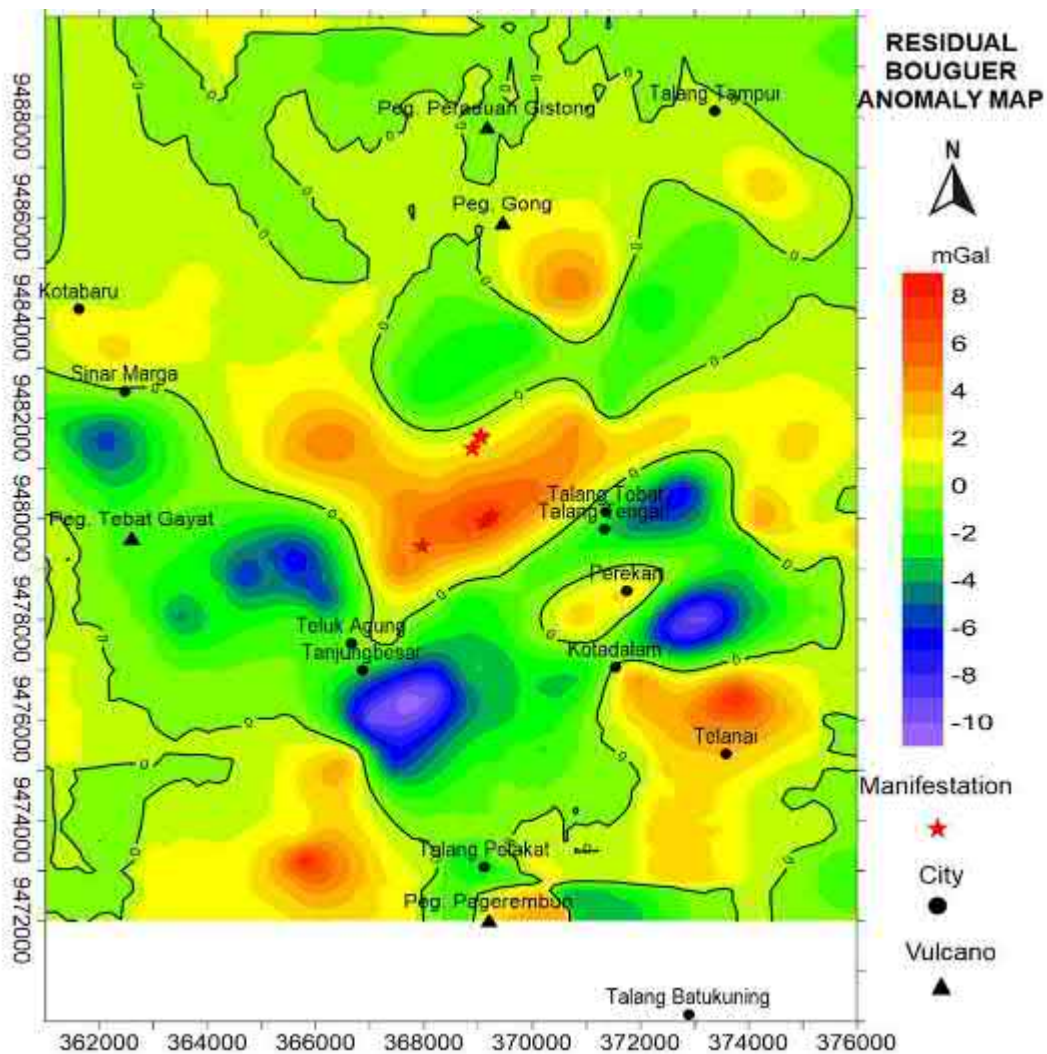


Figure 5. Residual Bouguer Anomaly Map of the Study Area

Second Vertical Derivative (SVD)

The Elkins operator generates the Second Vertical Derivative (SVD) map from the anomaly filter process (Elkins, 1951). Black contour lines indicate anomaly contours with a value of zero. Zero contours flanked by high and low anomaly values indicate the existence of a fault structure in the area. From the SVD map, a fault structure line can be drawn against the anomaly contour with a value of zero. In the Second Vertical Derivative principle, the value of zero indicates the value of a boundary, so it can be assumed that the area has several faults.

Second Vertical Derivatives analysis was performed on residual anomalies.

Figure 7 shows the SVD contour map of the study area. The filtered SVD contour map has anomaly values in the range of -8.5 to 5.5 mGal. The anomaly value is divided into three parts, namely low anomaly (-8.5 to -3.5 mGal), moderate gravity value (-2.5 to 2.5 mGal), and high gravity value (> 2.5 mGal). The residual SVD contour map shows high SVD values in orange to red and low SVD values in purple to blue. In the center of the study area, the zero anomaly contours are identified as the boundaries of the geothermal reservoir.

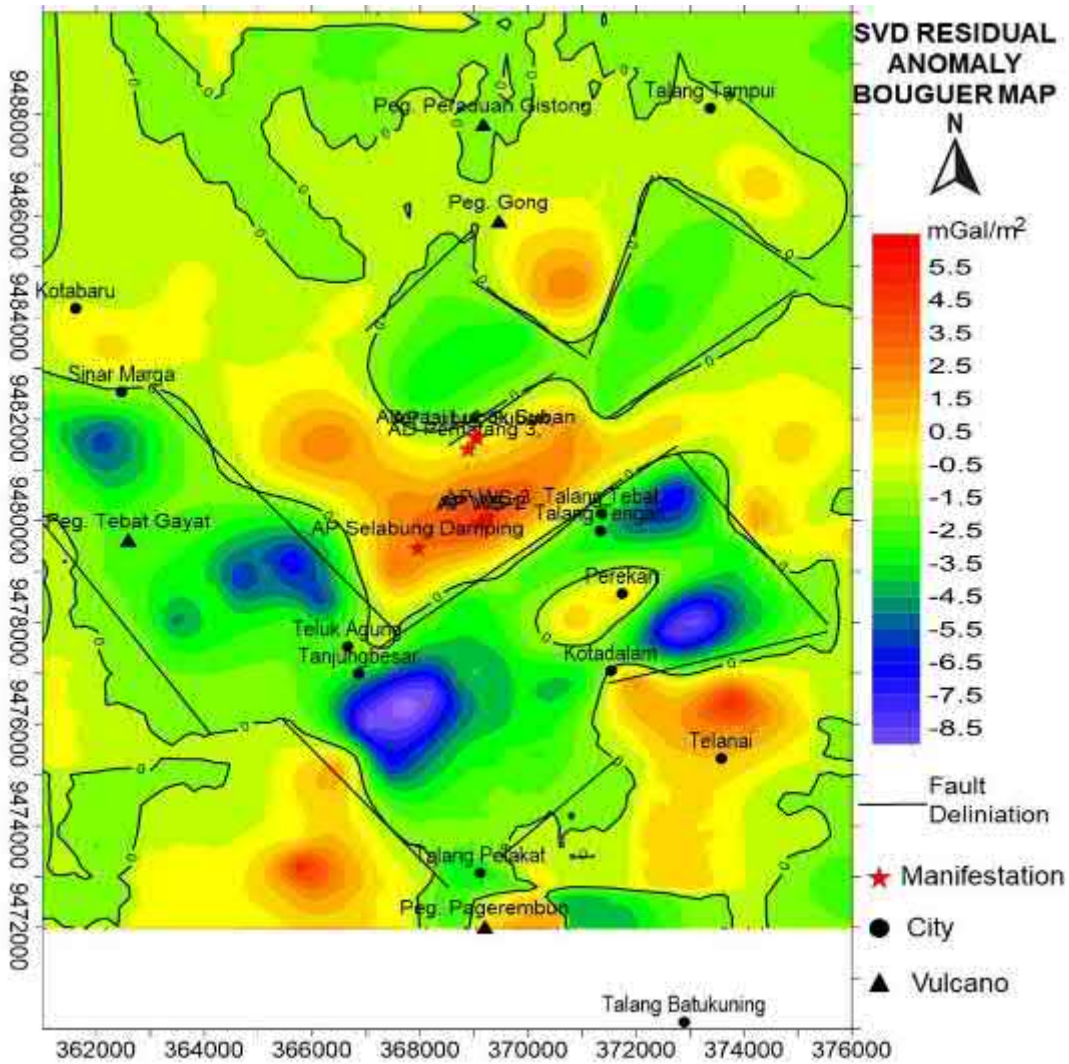


Figure 6. SVD Residual Anomaly Bouguer Map

2D Apparent Resistivity Distribution Model

Based on the results of an integrated investigation in 2011, line 1 consists of measurement points MTWS-42, MTWS-41, MTWS-25, MTWS-39, MTWS-03, MTWS-40, MTWS-08, MTWS-13, and MTWS-18, as shown in picture 7. Line 1 is southwest to northeast. The distribution of low resistivity value that spreads below the point MTWS-42 to MTWS-18 and leads to the northeast of the low is thought to be a rock cover layer.

The moderate resistivity value pattern distribution with a value range of 45–380 Ωm is considered a reservoir because it is located in two zones with high and low contrast resistivity. A reservoir is a volume of a permeable rock layer that stores heat and a circulation fluid that extracts heat.

The distribution of high resistivity value (>380 Ωm) is thought to be a basement experiencing intrusion. It is suspected to be one of the heat sources given the high resistivity value compared to other zones.

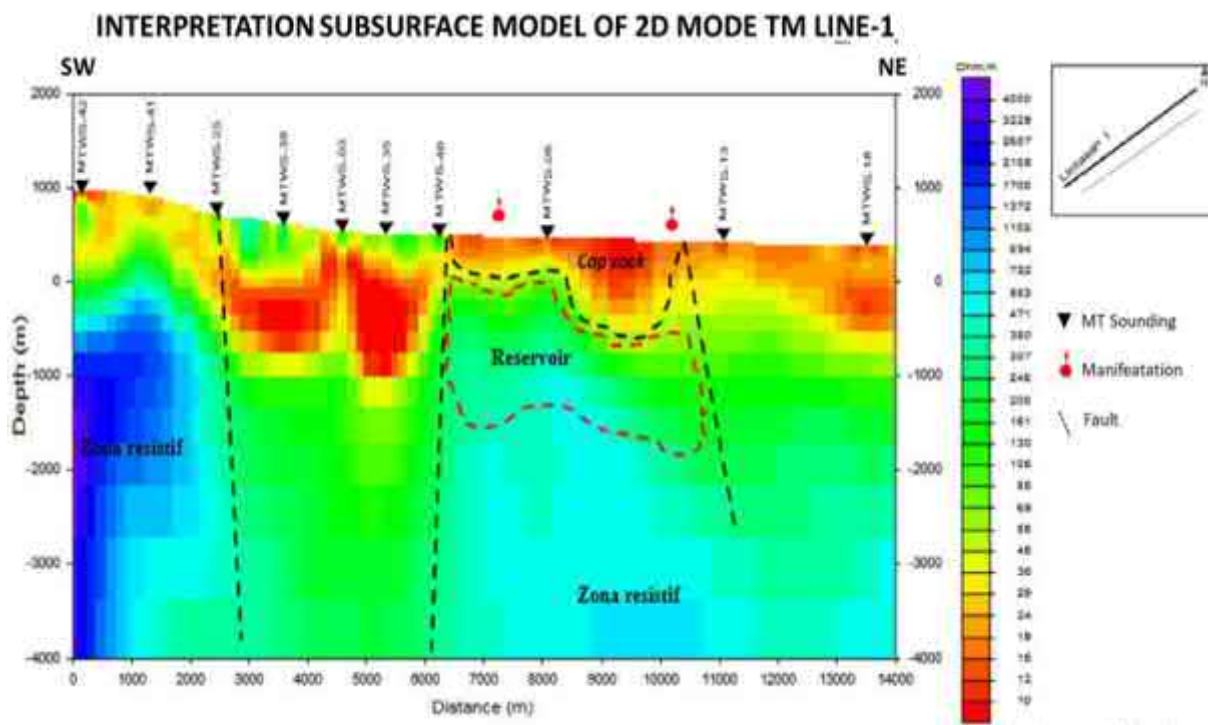


Figure 7. 2D Apparent Resistivity Model of Line 1(Geothermal Directorate, 2017)

Line 2 runs from southwest to northeast can be seen in figure 8. The low resistivity layer of 1-45 Ωm near the surface is considered the rock cover layer. Layers with

a moderate resistivity value (45-380 Ωm) are interpreted as reservoirs, while moderately high resistivity (380-4000 Ωm) is assumed as basement rocks.

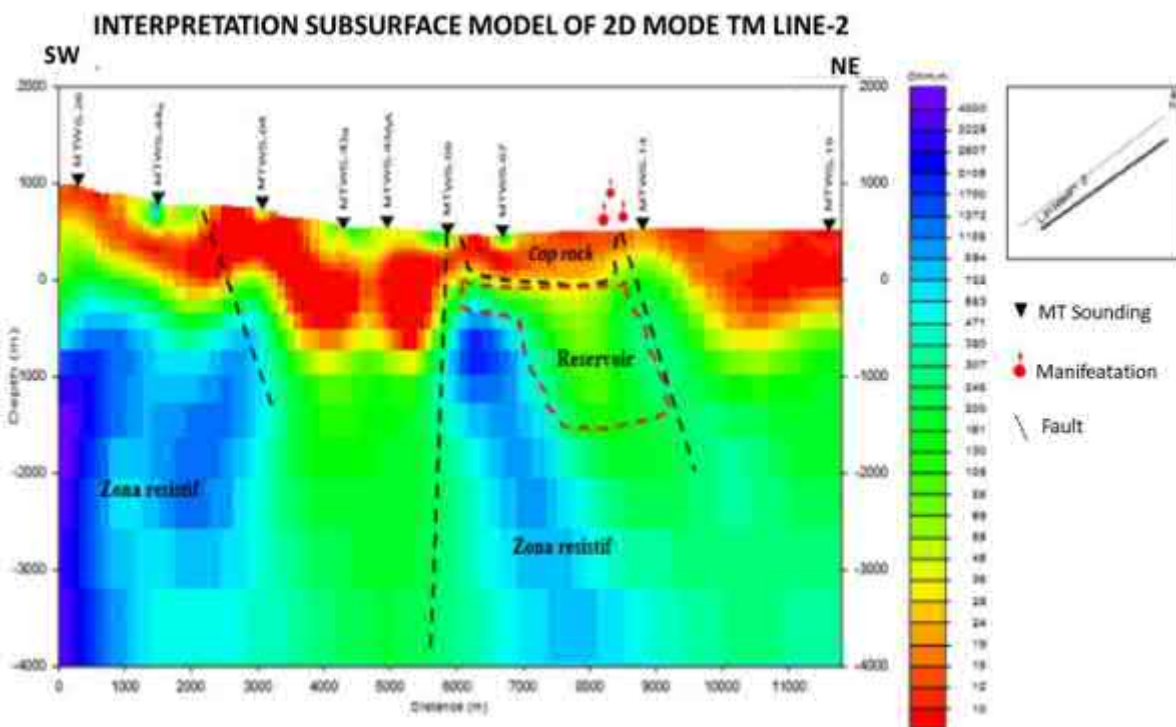


Figure 8. 2D Apparent Resistivity Model of Line 2 (Geothermal Directorate, 2017)

3D Modeling of Gravity

The residual Bouguer anomaly 3D inversion modeling was carried out using Grablox software. This inversion modeling uses a box-shaped object model approach arranged in a grid with specific dimensions, as input is the residual Bouguer anomaly grid data. After running the data process, the output in the form of a density value is obtained, then displayed in a 3D model as shown in Figure 9. The inversion of the residual anomaly (seen from the southwest) produces a density distribution as shown in Figure 8 with a value range of 2 to 2.9 g/cm³ with a maximum depth of 4000 meters. The distribution of low surface density values is in the north to a west area with a density value of 2 to 2.15 g/cm³. Low-density values

indicate the location of the reservoir. This area is located in the geological setting of old Miocene breccia, which was overwritten by Pleistocene Tebat Gayat lava. The distribution value of medium density ranges from 2.2 to 2.4 g/cm³. The density distribution is in the middle to the ends of the study area; the north to the west has a depth of 225 to 4,000 meters. Medium-density values are probably tertiary-aged sandstone sediment deposits. The distribution value of high-density ranges from 2.5 to 2.9 g/cm³. High-density values are spread at a depth of 225 to 4,000 meters; high-density values are of interest in gravity surveys because they indicate potential heat sources and possible hydrothermal silicate deposits (Sarkowi & Wibowo, 2021).

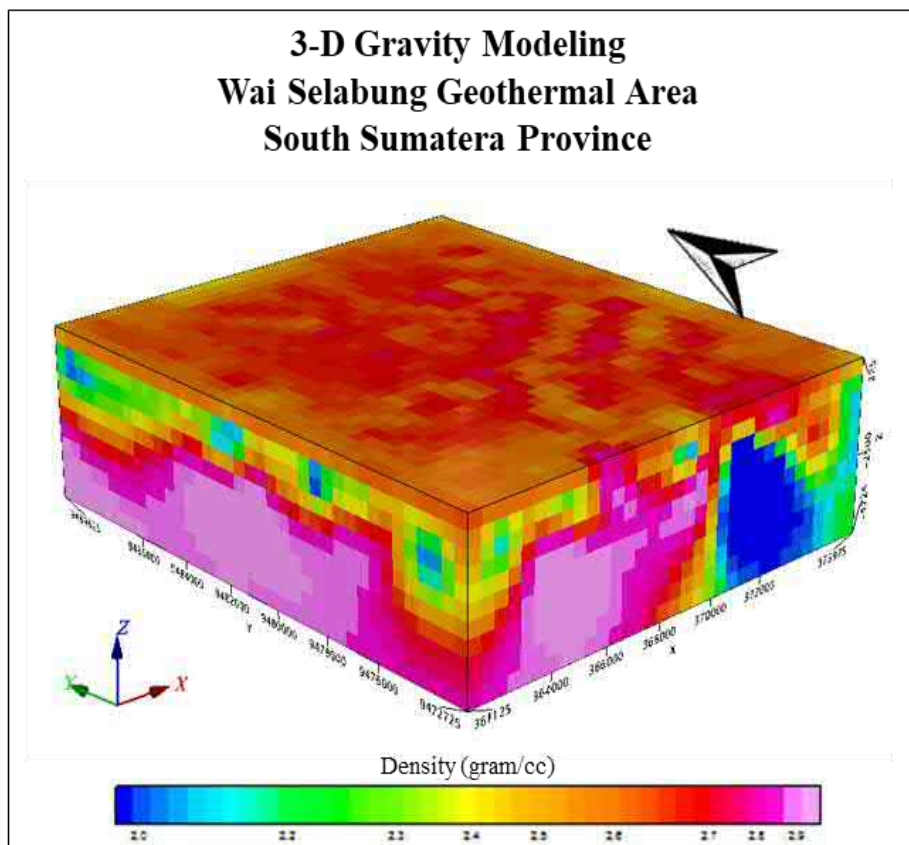


Figure 9. 3-Dimensional Model of the Research Area

The density distribution model resulting from the 3D inversion of the residual Bouguer anomaly above is then sliced to obtain a 2D cross-section of the trajectory.

Figure 10 shows the location of the reservoir in the study area based on the

analysis of the density distribution of the study area. The A-A' trajectory is northwest-southeast of the study area. The fault model is derived from a horizontal cross-section of order-1 (maximum or minimum) and order-2 (zero) gradient, while low density is assumed

to be a reservoir in the area. The Pematanggong fault influences the fault in this trajectory with a southwest-northeast

trending fault pattern and the Peraan fault with a north-south trending fault pattern.

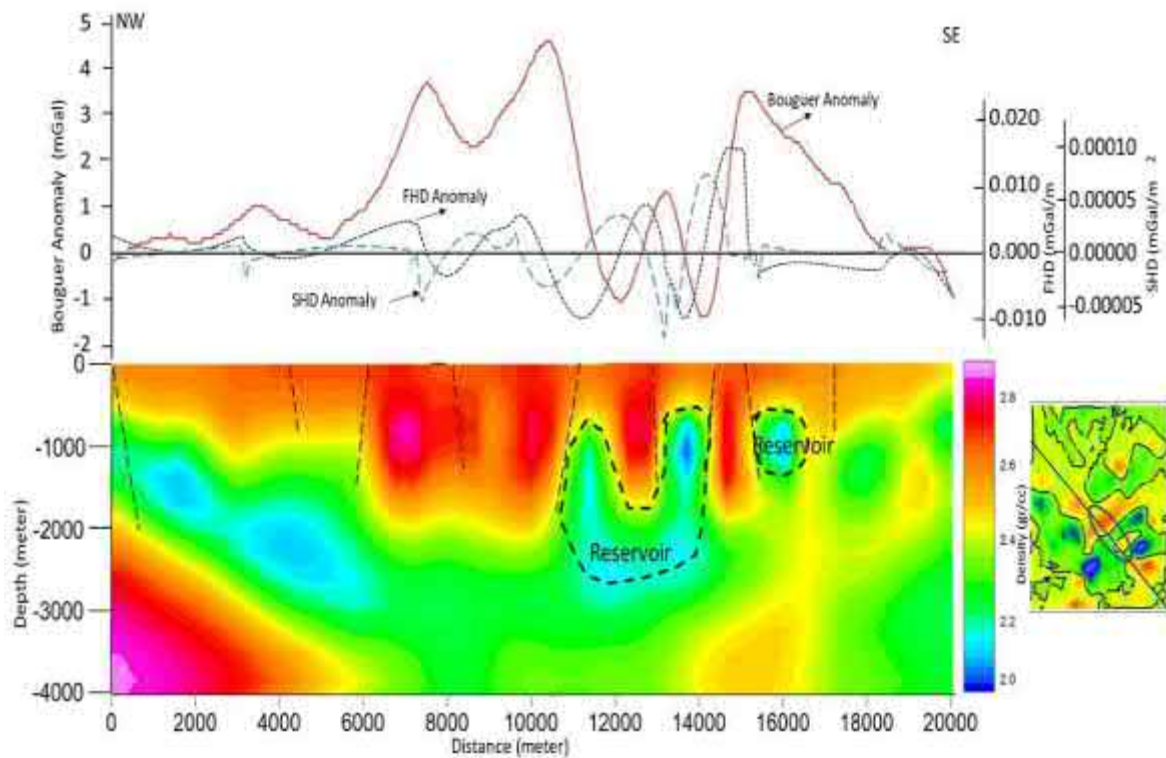


Figure 10. Density Distribution Cross-Sectional Model Trajectory A-A 'Section Results from the 3D Model Density Inversion Modeling

Figure 11 shows the location of the reservoir in the study area based on the analysis of the density distribution of the study area, which is correlated with the 1 MT trajectory. The B-B's trajectory is in the southwest-northeast direction of the study area. The fault model is derived from the cross-section of the first-order horizontal gradient (maximum or minimum) and the second-order gradient (zero). At the same time, the low density near the manifestation is assumed to be the reservoir in the area. The reservoir's existence is also supported by the

MT model, which shows the existence of a reservoir in the area limited by faults. The fault in this path is supported by the emergence of hot springs. The Pematanggong fault influences the faults in this trajectory with a southwest-northeast trending fault pattern, the Wai Selabung fault and the Kotadalam fault, which control the manifestation of the Selabung Damping AP, and the Gistong fault with a southwest-northeast pattern which controls the occurrence of the Lubuk Suban AP manifestation.

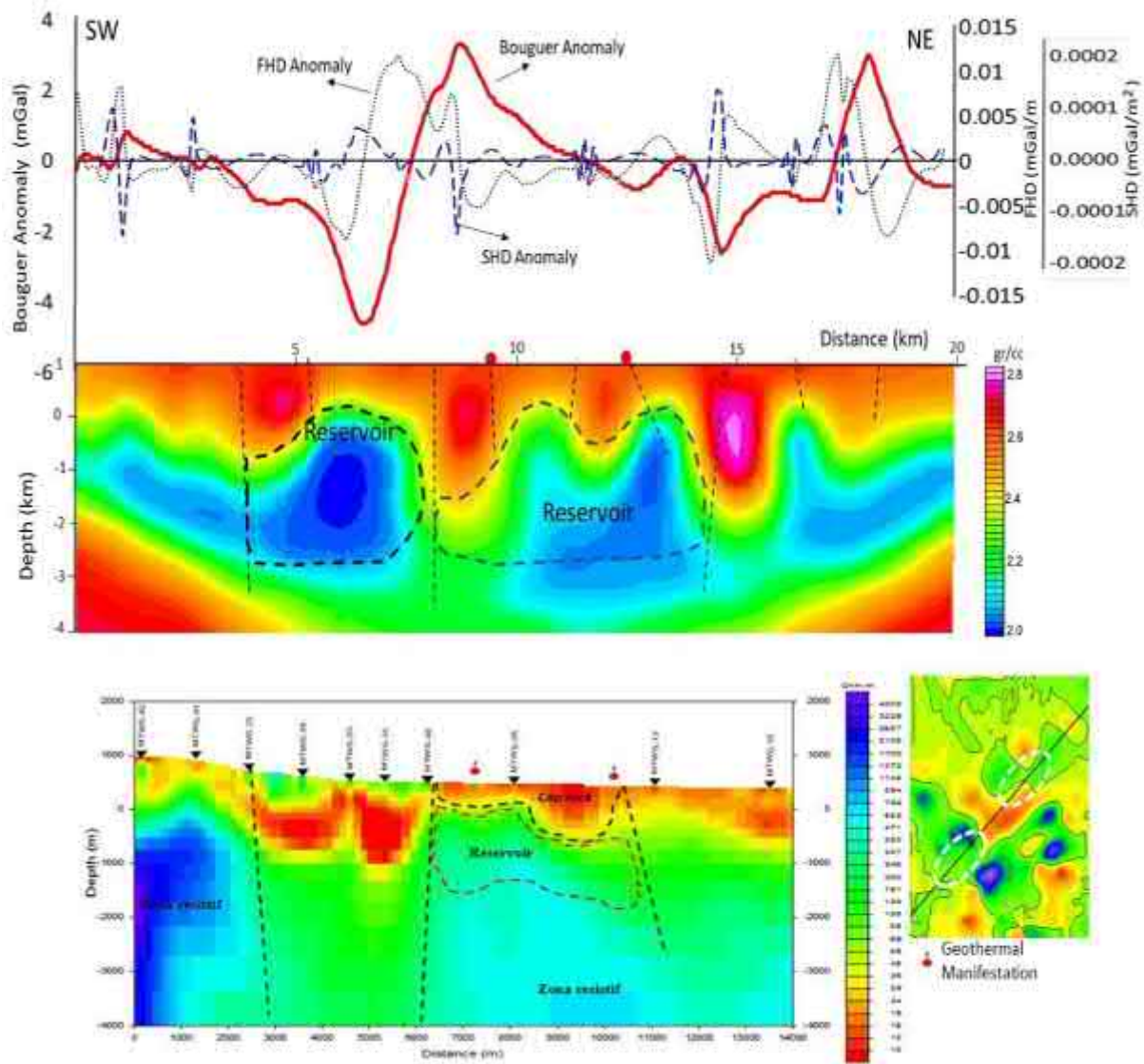


Figure 11. Density Distribution Cross-Sectional Model Trajectory B-B 'Section Results from the 3D Model Density Inversion Modeling

Figure 12 shows the reservoir based on the density distribution analysis, which is correlated with the 2 MT model. The C-C' cross-section is southwest-northeast of the study area. The fault model is derived from the cross-section of the 1st order horizontal gradient (maximum or minimum) and the 2nd order gradient (zero). At the same time, the low density near the manifestation is assumed to be the reservoir in the area. The reservoir's existence is also supported by the MT model, which shows the existence of a

reservoir in the area limited by faults. The fault is one component of the geothermal system that functions as a fluid transport route from the reservoir. The number of fault zones will increase the amount of fluid that comes to the surface. This fault is reinforced by the presence of hot springs near track 3. The Akarjangkang fault influences the fault in this line with a north-south trending fault pattern that controls the emergence of the manifestations of AP Wai Selabung 1 and AP Wai Selabung 2.

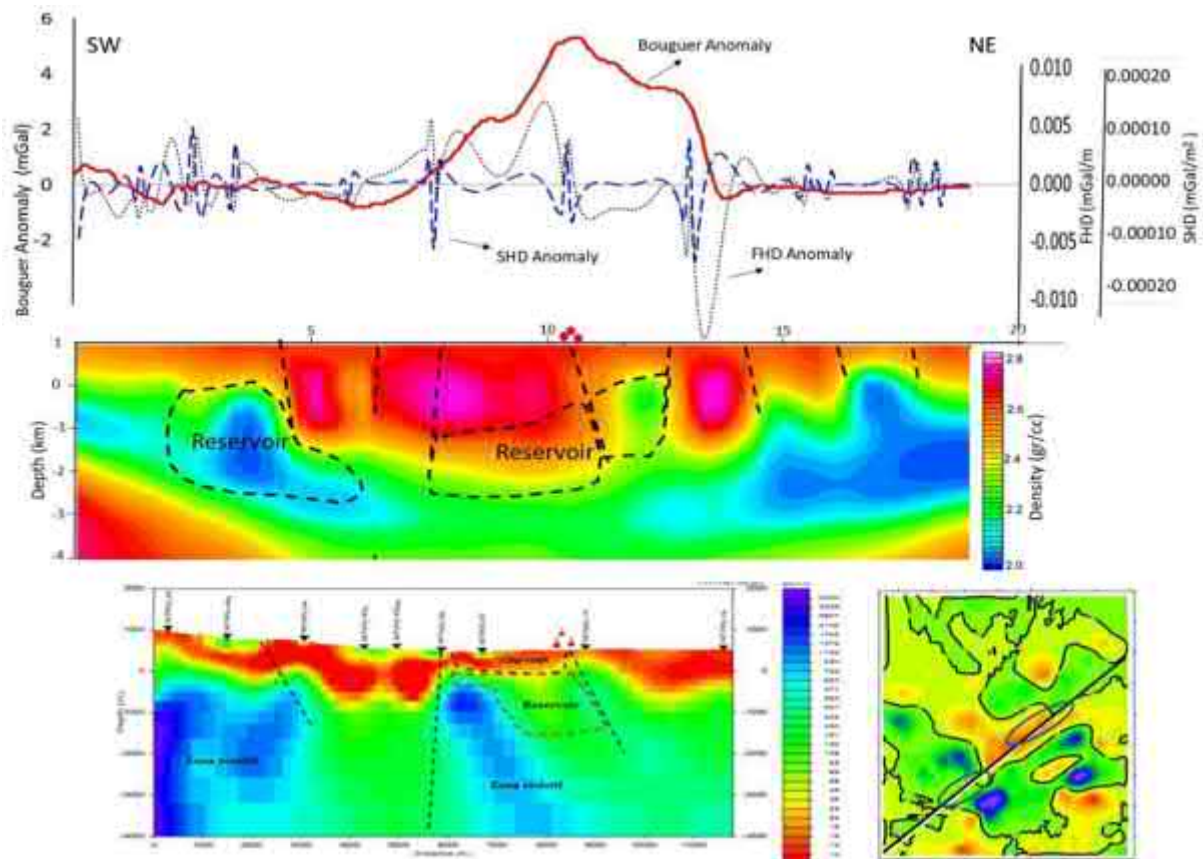


Figure 12. Density Distribution Cross-Sectional Model Trajectory C-C' Section Results from the 3D Model Density Inversion Modeling

Figure 13 shows the location of the reservoir in the study area based on the analysis of the density distribution of the study area. The D-D' trajectory is in the southwest-northeast direction of the study

area. The fault model is derived from a horizontal cross-section of order-1 (maximum or minimum) and order-2 (zero) gradient, while low density is assumed to be a reservoir in the area.

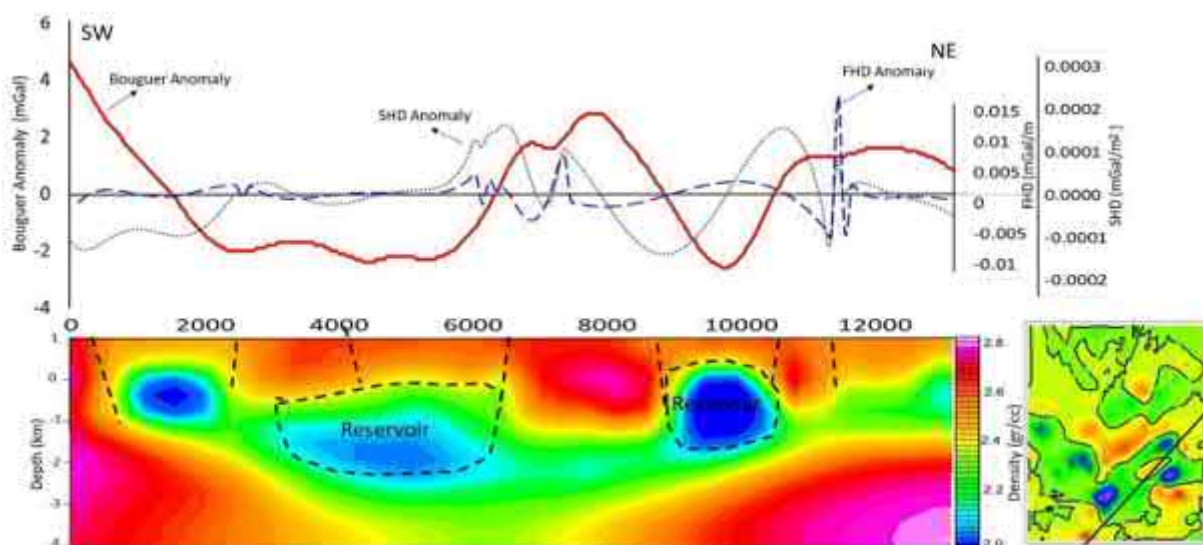


Figure 13. Density Distribution Cross-Sectional Model Trajectory D-D' Section Results from the 3D Model Density Inversion Modeling

Figure 14 shows the location of the reservoir in the study area based on the analysis of the density distribution of the study area. The E-E' trajectory is in the southwest-northeast direction of the study

area. The fault model is derived from a horizontal cross-section of order-1 (maximum or minimum) and order-2 (zero) gradient, while low density is assumed to be a reservoir in the area.

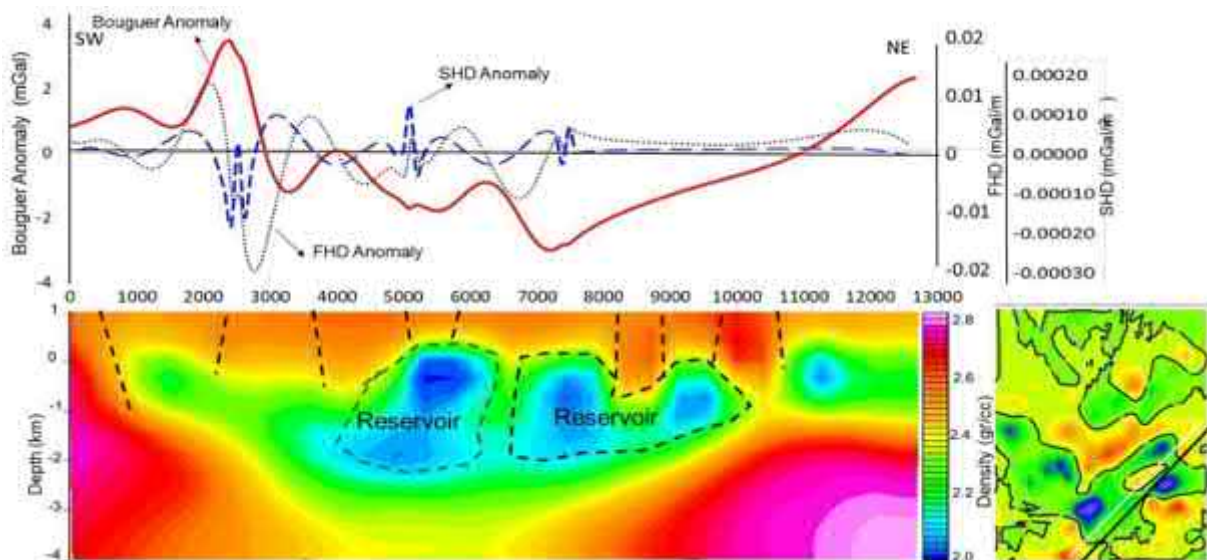


Figure 14. Density Distribution Cross-Sectional Model Trajectory E-E' Section Results from the 3D Model Density Inversion Modeling

Geochemical Analysis

Based on the results of an integrated investigation in 2011, the characteristics and types of hot water are based on the Cl-SO₄-HCO₃ triangle, which refers to Giggenbach (1988) with the type of chloride bicarbonate or chloride sulfate and type bicarbonate.

In determining the temperature of the geothermal reservoir in the research area, the Na-K-Mg triangle is used. Wai Selabung 1 hot spring and Wai Selabung 2 hot springs are in partial equilibrium and are thought to originate from the reservoir with a temperature of 160°C to 180°C. The partial equilibrium zone indicates that the fluid reaction with the reservoir rock has reached partial equilibrium. The other hot springs samples are in the immature water zone, indicating that the hot water has been mixed with cold water on the surface in a high proportion.

To determine the origin of the recharge reservoir water using the stable isotopes D-deuterium and ¹⁸O. The plotting results of Wai Selabung 2 hot springs and Wai Selabung 3 hot springs are close to the

meteoric water line, indicating that the hot springs are influenced by surface water. Meanwhile, the other hot springs, namely Wai Selabung 1, Lubuk Suban, Selabung Damping, Arumatai, and Kota Batu, show oxygen enrichment. So, that the position on the graph is to the right of the meteoric water line. An indication of the formation of Hot springs is related to the interaction between the hot fluid in the geothermal system and the rock that causes ¹⁸O enrichment. This indication shows that Wai Selabung 1 Hot Water, Lubuk Suban Hot spring, Selabung Damping Hot spring, Arumatai Hot Spring, and Kota Batu Hot spring are more influenced by magmatic fluids.

In the 2011 integrated investigation, the bottom borehole temperature, surface temperature gradient, and surface heat flow were also carried out. The results obtained that the distribution of the anomaly zones of the three measurements is around the manifestation of the Wai Selabung hot spring to the Lubuk Suban hot spring.

Geothermal Reservoir Prospect Area Analysis

Previous research by the Center for Geological Resources in 2011 showed that the prospect area is only in the southwest (Kusnadi et al., 2011). The analysis of gravity, MT, geochemistry, and geology of the study area is depicted in a compilation map showing the reservoir prospect area of the study area, as shown in Figure 15. The prospect area for the Wai Selabung geothermal reservoir is estimated to be located in old lava rock and sandstone sedimentary layers under the Ranau

Pyroclastic layer, located around the Great Bay. This area is influenced by the Teluk Agung fault, which has a northwest-southeast trend. The next reservoir prospect area is in the sandstone sediment layer under the Sapatuhu Pyroclastic layer and the Peraan Lava rock around Perekani. The north-south Peraan fault influences this area. Moreover, the last reservoir prospect area is in the Akarjangkang Lava Rock and the shale sandstone sediment layer around Talang Tebat. This area is influenced by the Pematanggong fault and the Gistong fault, which are trending southwest-northeast.

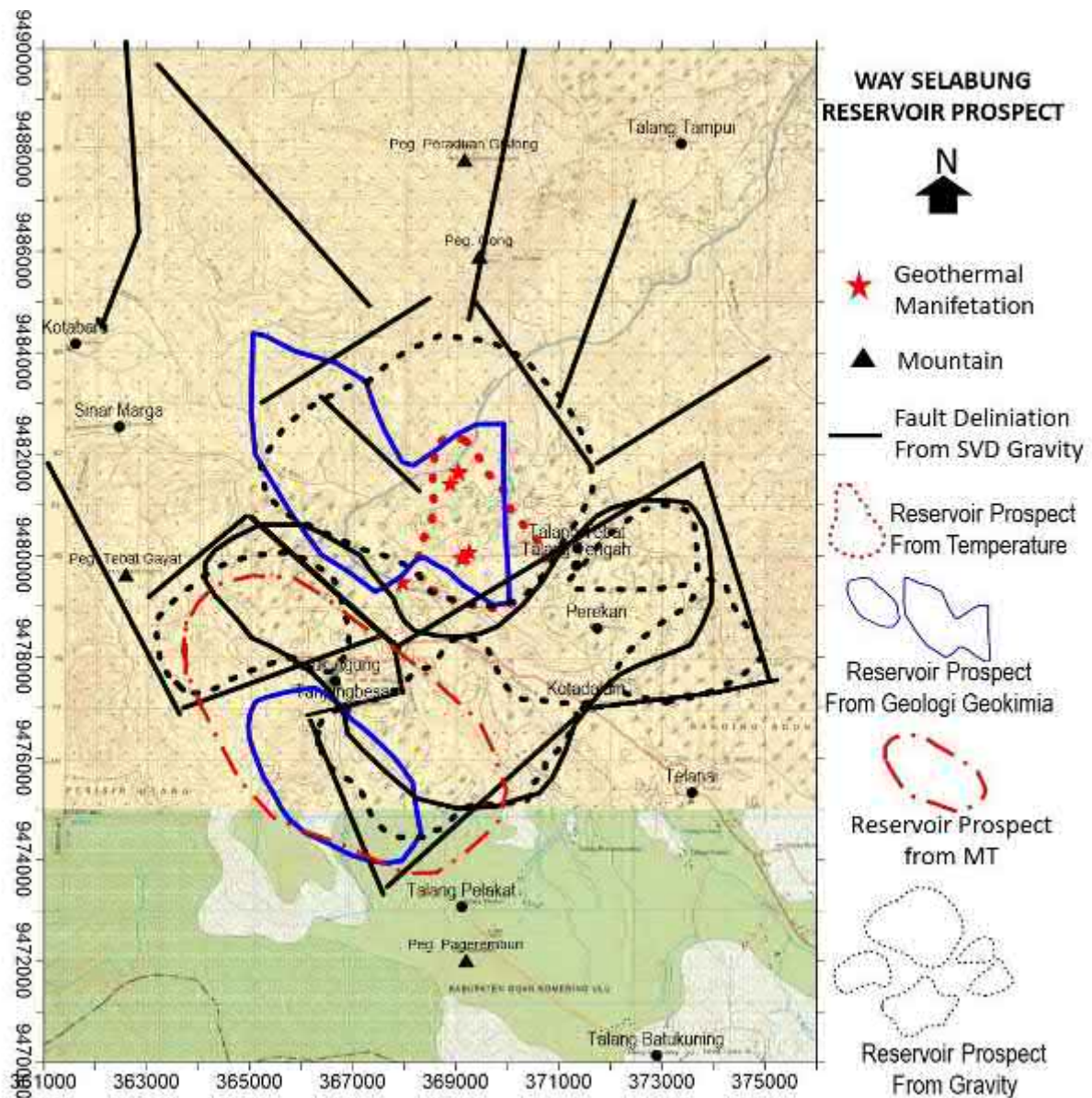


Figure 15. Map Compilation of Geology, Geochemistry, Gravity, and MT of Wai Selabung Geothermal Area

CONCLUSION AND SUGGESTION

The conclusions of this research can be drawn: the prospect area for the Wai Selabung geothermal reservoir based on the Bouguer anomaly inversion modeling, which is correlated with geological data, geochemical data, and magnetotelluric data, is estimated to be located around Teluk Agung, Perekan, and Talang Tebat. This reservoir prospect area can be exploited to meet energy needs in Indonesia.

It is necessary to conduct exploration drilling studies to support the results of subsurface research and to calculate the potential of geothermal energy in the research area so that a feasibility study can be carried out for the development of geothermal fields.

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REFERENCES

- Abdelrahman, E. M., & El-Araby, T. M. (1996). Shape and depth solutions from moving average residual gravity anomalies. *Journal of Applied Geophysics*, 36, 89–95.
- Andari, W., Karyanto, & Kurniawan, R. (2019). Identifikasi batas sub-cekungan hidrokarbon menggunakan analisis SHD (second horizontal derivative) dan SVD (second vertical derivative) berdasarkan korelasi data gayaberas dan seismik. *Jurnal Geofisika Eksplorasi*, 5(1), 60–74. <https://doi.org/10.23960/jge.v>
- Aulia, Y. F., Karyanto, & Rustadi. (2019). Pendugaan patahan daerah “Y” berdasarkan anomali gayaberas dengan analisis derivative. *Jurnal Geofisika Eksplorasi*, 5(1), 75–88. <https://doi.org/10.23960/jge.v>
- Azizi, H. A., Haryanto, A. D., Hutabarat, J., Muslim, D., Gentana, D., Geologi, F. T., Padjadjaran, U., & Utara, P. S. (2020). Pola aliran panas berdasarkan analisis fluida dan mineral ubahan di daerah panas bumi sorik marapi, kabupaten mandailing natal, provinsi sumatra utara. *Buletin Sumber Daya Geologi*, 15(3), 202-218.
- Bjornsson, G., & Bodvarsson, G. (1990). A survey of geothermal reservoir properties. *Geothermics*, 19(1), 17–27.
- Blakely, R. J. (1996). *Potential theory in gravity and magnetic applications*. Cambridge University Press.
- Brehme, M., Deon, F., Haase, C., Wiegand, B., Kamah, Y., Sauter, M., & Regenspurg, M. (2016). Fault controlled geochemical properties in Lahendong geothermal reservoir Indonesia. *Grundwasser – Zeitschrift Der Fachsektion Hydrogeologie*, 21, 29–41. <https://doi.org/10.1007/s00767-015-0313-9>
- Daud, Y., Rosid, M. S., Pati, G. P., Maulana, M. R., & Khoiroh, M. (2018). Imaging structural control of geothermal reservoir using remote sensing and gravity data analysis in Blawan-Ijen, East Java, Indonesia. *AIP Conference Proceedings*, 2023, 1–6. <https://doi.org/10.1063/1.5064260>
- Dewi, I. K., Puspitasari, F., Z, N. M., & Martha, A. (2020). 3D data gravity modeling for identification of the formation structure of the hydrocarbon basin in the bajubang region, jambi province. *Jurnal Geofisika Eksplorasi*, 6(3), 216–227. <https://doi.org/10.23960/jge.v6i3.103>
- Elkins, T. A. (1951). The second derivative method of gravity interpretation. *SEG*, 15, 59–81.
- Giggenbach, W. F. (1988). Geothermal solute equilibria. Derivation of Na-K-Mg-Ca geothermometers. *Geochemica et Cosmochimica Acta*, 52(12), 2749–2765.
- Hadi, M. N., Munandar, A., Kusnadi, D., Zarkasyi, A., Suryakusuma, D., & Sugianto, A. (2011). Survei terpadu geologi, geokimia, dan geofisika daerah panas bumi wai selabung, kabupaten

- oku selatan, provinsi sumatera selatan. *Prosiding Hasil Kegiatan Pusat Sumber Daya Geologi Tahun, 1-14*.
- Haerudin, N., Wahyudi, W. S., Sarkowi, & D. R. (2014). Magnetic and gravity modeling to determine reservoir depth and prospect area at rajabasa lampung. *International Conference on Mathematics, Science, and Education 2014*, 37–44.
- Haerudin, Nandi, Pardede, V. J., & Rasimeng, S. (2009). Analisis reservoir daerah potensi panas bumi gunung rajabasa kalianda dengan metode tahanan jenis dan geotermometer. *Jurnal Ilmu Dasar*, 10(2), 141–146.
- Hafidah, A. D., Daud, Y., & Usman, A. (2019). Reservoir identification based on gravity method at "AUN" geothermal field. *E3S Web of Conferences*, 125(ICENIS 2019), 1–6.
- Ilmi, I., Syafri, I., Haryanto, A. D., & Zarkasyi, A. (2020). Pemodelan inversi 2-d menggunakan data magnetotellurik daerah panas bumi way selabung, kabupaten ogan komering ulu selatan, provinsi sumatera selatan. *Buletin Sumber Daya Geologi*, 15(1), 61-72.
- Ilmi, N. (2019). Pemodelan 3D Struktur bawah permukaan gunungapi agung provinsi bali menggunakan metode gaya berat. *Jurnal Geofisika Eksplorasi*, 5(2), 68-75.
- Imran, Brandon, P., Fernanda, E., & Sudrazat, S. D. (2021). Pengolahan data landsat dan gravitasi sebagai indikasi panas bumi daerah rana kulan, NTT. *Jurnal Geofisika Eksplorasi*, 07(01), 41–51.
- Iqbal, M., Juliarka, B. R., Ashuri, W., & Farishi, B. Al. (2019). Hydrogeochemistry of natar and cisarua hot springs in south lampung, indonesia. *Journal of Geoscience, Engineering, Environment, and Technology*, 4(3), 178–185. <https://doi.org/10.25299/jgeet.2019.4.3.4070>
- Irawati, S. M., Luthfian, A., & Laesanpura, A. (2021). Subsurface structure of baturagung escarpment revealed through three-dimensional gravity inversion. *Jurnal Geofisika Eksplorasi*, 7(1), 17–29. <https://doi.org/10.23960/jge.v7i1.125>
- Iswahyudi, S., Attabik, L. W., Setijadi, R., & Raharjo, S. A. (2019). Perhitungan temperatur reservoir panas bumi mata air panas daerah bantarkawung menggunakan metode geotermometer Na-K dan entalpi-silika. *Jurnal Geosaintek*, 5(1), 7-12. <https://doi.org/10.12962/j25023659.v5i1.5055>
- Kusnadi, D., Risdianto, D., Munandar, A., & Dahlan. (2011). Geologi dan geokimia daerah panas bumi wai selabung kabupaten oku selatan, sumatera selatan. *Prosiding Hasil Kegiatan Pusat Sumber Daya Geologi Tahun*, 1–14.
- Maulana, Y., Haryanto, A. D., Ismawan, & Kusnadi, D. (2019). Penentuan tipe fluida serta tingkat permeabilitas sistem panas bumi daerah tinigi, toli-toli, sulawesi tengah. *Padjadjaran Geoscience Journal*, 3(6), 459–464. <http://journal.unpad.ac.id/geoscience/article/view/26335>
- Motta, J. G., Souza Filho, C. R. d., Carranza, E. J. M., & Braitenberg, C. (2019). Archean crust and metallogenic zones in the Amazonian Craton sensed by satellite gravity data. *Scientific Reports*, 9(1), 1–10. <https://doi.org/10.1038/s41598-019-39171-9>.
- Nasution, A., Hassan, A.R., Kasbani, Bakrun, & Sulaeman, B. (1999). *Penyelidikan potensi panas bumi di kabupaten ogan komering ulu (OKU) Sumatera Selatan*. Pemaparan Hasil Kegiatan Lapangan DIK-S Batubara, DSM, Bandung.
- Permada, R., & Putra, A. (2017). Estimasi karakteristik reservoir panas bumi dari sumber mata air panas di kecamatan pauh duo, kabupaten solok selatan. *Jurnal Fisika Unand*, 6(1), 39–46.

- Permatasari, A. W., Santoso, D., & Setianingsih. (2019). Subsurface modeling based on gravity data of rendingan-ulubelu- subsurface (RUW) geothermal field, Lampung, Indonesia. *IOP Conf. Series: Earth and Environmental Science*, 318, 1–7. <https://doi.org/10.1088/1755-1315/318/1/012002>
- Polii, J., & Rampengan, A. M. (2020). Analisa geokimia fluida manifestasi permukaan di daerah panas bumi lahendong. *Fullerene Journal of Chemistry*, 5(1), 45–48. <https://doi.org/10.37033/fjc.v5i1.148>
- Prasetyo, R., Laksmeningpuri, N., & Pratikno, B. (2017). Karakterisasi isotop dan geokimia area panas bumi danau toba, sumatera utara. *Jurnal Ilmiah Aplikasi Isotop Dan Radiasi*, 13(2), 79–86. <https://doi.org/10.17146/jair.2017.13.2.3508>
- Putri, D. A. B., & Harianja, A. K. (2021). Identifikasi prospek panas bumi radiogenik menggunakan landsat-8 dan gravitasi di daerah permis. *Jurnal Geofisika Eksplorasi*, 07(01), 52–70.
- Putriyana, L., & Soekarno, H. (2020). Numerical modelling of geothermal reservoir in gunung talang, west sumatera, indonesia. *Jurnal Geologi Riset Dan Pertambangan*, 30(1), 11–20. <https://doi.org/10.14203/risetgeotam2020.v30.1058>
- Rizkiani, D. N., & Rustadi. (2019). Interpretasi sistem panas bumi suwawa berdasarkan data gaya berat. *Jurnal Geofisika Eksplorasi*, 5(2), 130–140. <https://doi.org/10.23960/jge.v>
- Sarkowi, M. & Wibowo, R.C. (2021). Reservoir identification of Bac-Man geothermal field based on gravity anomaly analysis and modeling. *Journal of Applied Science and Engineering*, 25(2), 329–338. doi: [http://dx.doi.org/10.6180/jase.202204_25\(2\).0009](http://dx.doi.org/10.6180/jase.202204_25(2).0009)
- Setiadi, I., Setyanta, B., & Widijono, B. S. (2010). Delineasi cekungan sedimen sumatra selatan berdasarkan analisis data gaya berat. *Jurnal Sumber Daya Geologi*, 20(2), 93–106.
- She, Y., & Fu, G. (2020). Uplift mechanism of the highest mountains at eastern himalayan syntaxis revealed by in situ dense gravimetry. *Geophysical Research Letters*, 47(22), 0–3. <https://doi.org/10.1029/2020GL091208>
- Sihombing, R. B., & Rustadi. (2018). Pemodelan dan analisa struktur bawah permukaan daerah prospek panasbumi kepahiang berdasarkan metode gayaberat. *Jurnal Geofisika Eksplorasi*, 4(2), 159–172. <https://doi.org/10.23960/jge.v4i2.14>
- Sulistyo, A., & Y, D. (2019). Delineating geothermal system through 3D geomagnetic and gravity data inversion on Blawan Ijen geothermal area , East Java. *IOP Conference Series: Earth and Environmental Science*, 538, 1–8. <https://doi.org/10.1088/1755-1315/538/1/012061>
- Sumintadireja, P., Dahrin, D., & Grandis, H. (2018). A note on the use of the second vertical derivative (SVD) of gravity data with reference to indonesian cases. *Journal of Engineering and Technological Sciences*, 50(1), 127–139. <https://doi.org/10.5614/j.eng.technol.sc.i.2018.50.1.9>
- Tripp, A., Moore, J., Ussher, G., & Mcculloch, J. (2002). Gravity modeling of the karaha - telaga bodas geothermal system, indonesia. *PROCEEDINGS, Twenty-Seventh Workshop on Geothermal Reservoir Engineering, SGP-TR-171*, 28–30.
- Utami, Z. D., & Putra, A. (2018). Penentuan karakteristik fluida dan estimasi temperatur reservoir panas bumi di sekitar gunung talang. *Jurnal Fisika Unand*, 7(2), 130–137.
- Van Camp, M., de Viron, O., Watlet, A., Meurers, B., Francis, O., & Caudron, C. (2017). Geophysics from terrestrial time-variable gravity measurements.

- Reviews of Geophysics*, 55(4), 938–992.
<https://doi.org/10.1002/2017RG000566>
- Wahyu, S., Daud, Y., Rahadinata, T., & Ningrum, S. S. (2019). Identifikasi potensi sistem panasbumi berdasarkan korelasi data geologi dan data gravitasi dengan menggunakan teknik filtering first horizontal derivative (FHD) - second horizontal derivatif (SVD). *Jurnal Teknologi*, 7(1), 40–53.
- Widiatmoko, F. R. (2019). Pendekatan analisa geokimia dengan multivariate analysis untuk mengetahui tipe mata air panas: studi kasus lapangan panas bumi mapos, nusa tenggara timur. *Jurnal IPTEK*, 23(2), 71–78.
<https://doi.org/10.31284/j.iptek.2019.v23i2>
- Witter, J. B., Siler, D. L., Faulds, J. E., & Hinz, N. H. (2016). 3D geophysical inversion modeling of gravity data to test the 3D geologic model of the Bradys geothermal area , Nevada , USA. *Geothermal Energy*, 4(14), 1–21.
<https://doi.org/10.1186/s40517-016-0056-6>
- Witter, J. B., Stelling, P., Knapp, P., & Hinz, N. H. (2016). 3D geophysical inversion modelling of gravity data as a subsurface geothermal exploration tool with an example from akutan (alaska , USA). *Geothermal Resources Council Transactions*, 40, 647–658.
- Yoni, D. R., Haryanto, A. D., Sapari, M., Hadian, D., & Atmaja, R. W. (2018). Karakterisasi geokimia fluida untuk menentukan sistem panasbumi dan potensi scaling lapangan patuha, jawa barat. *Padjajaran Geoscience Journal*, 2(4), 312–320.
- Zhu, Z., Lei, X., Xu, N., Shao, D., Jiang, X., & Wu, X. (2020). Integration of 3D geological modeling and geothermal field analysis for the evaluation of geothermal reserves in the northwest of beijing. *Water*, 12(638), 1–16.